

# Radar sensing of surface waves: An evaluation of waveheight estimates

Tyson Hilmer

Advisors:

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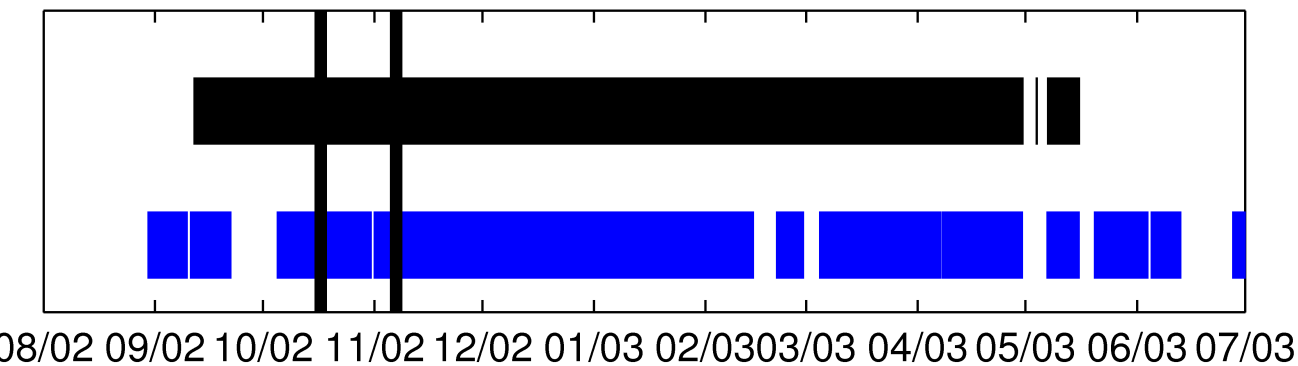




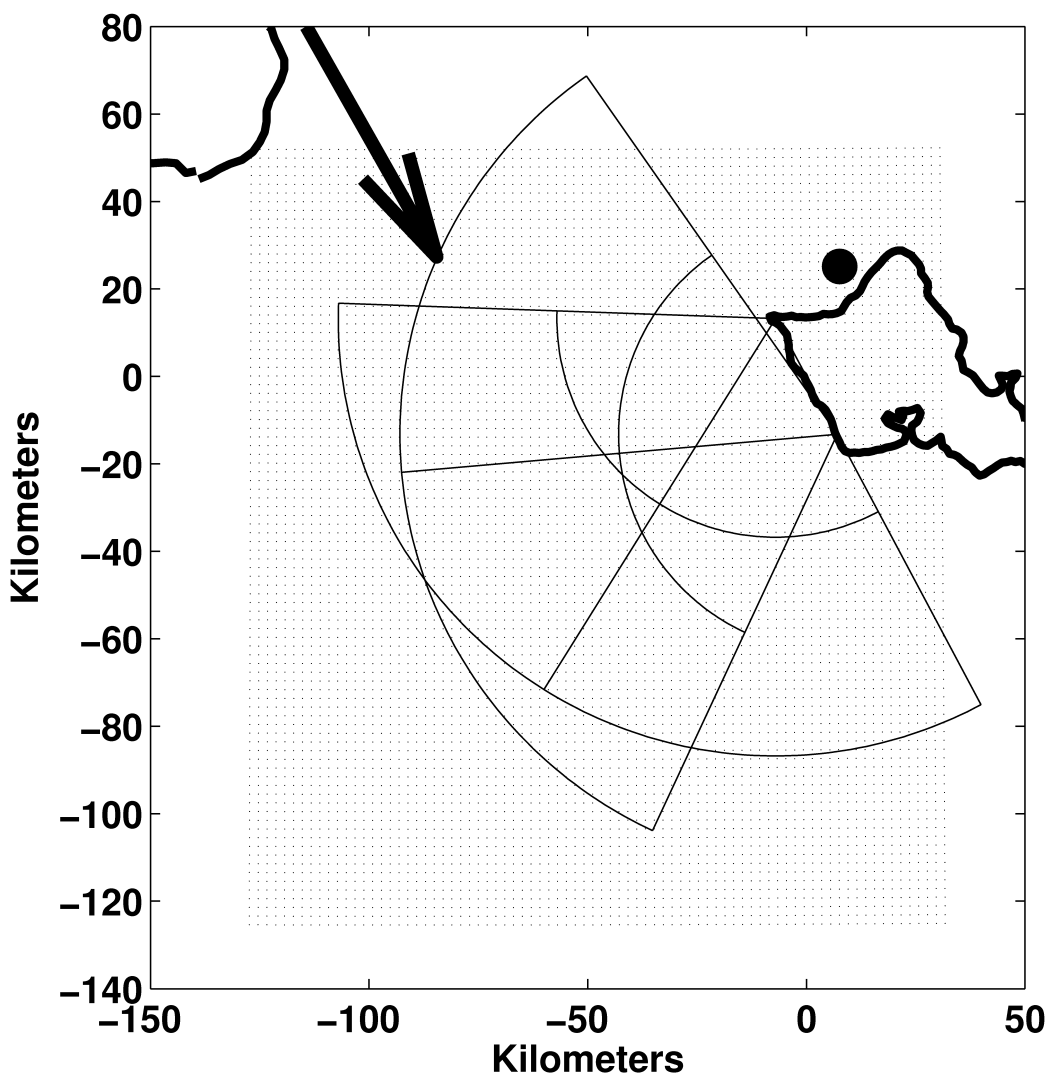
**Kaena**







**17-Oct to 06-Nov 2002, 21 days**

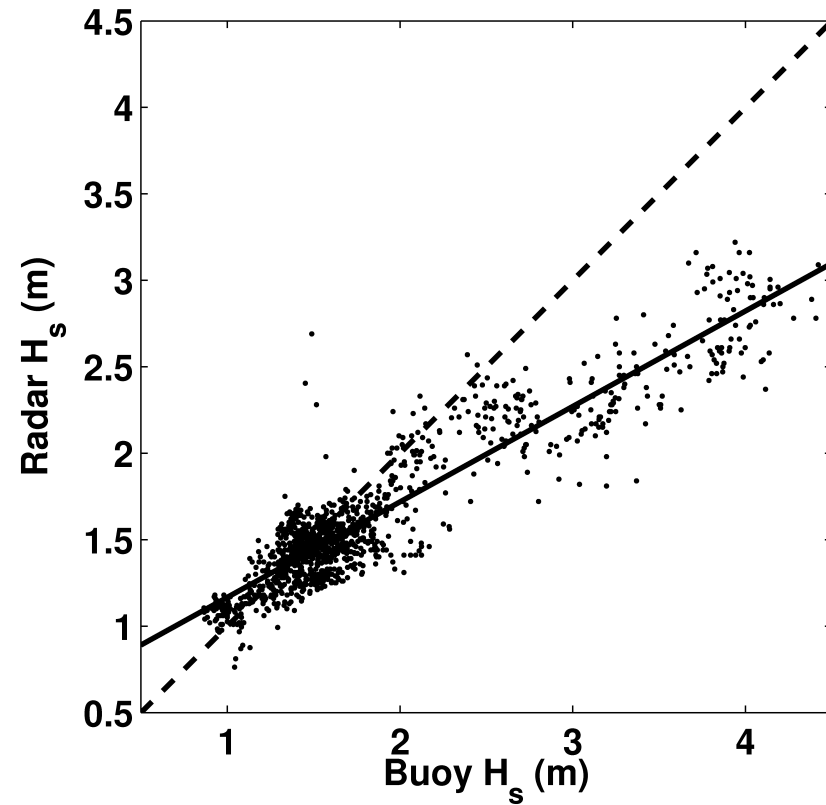
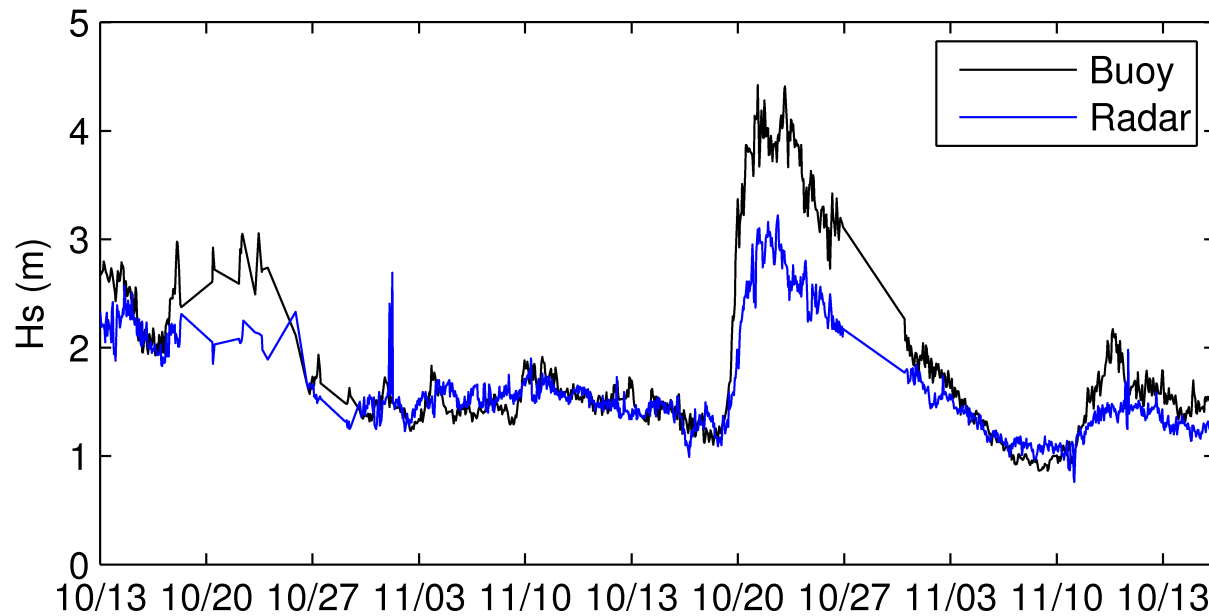


- 16 MHz transmit frequency
- 100 km nominal range
- 7.2 deg angular resolution
- 1.5 km range resolution

Bragg scattering ocean waves:

- 9.38 m wavelength
- 2.45 s period
- 0.41 Hz frequency

# Significant waveheight (Hs): Waimea Buoy & WERA



**Correlation  $r = 0.93$**   
**RMS error = 0.43 m**  
**RSE error = 0.16 %**



Table 1.1: Comparison studies of radar to buoys

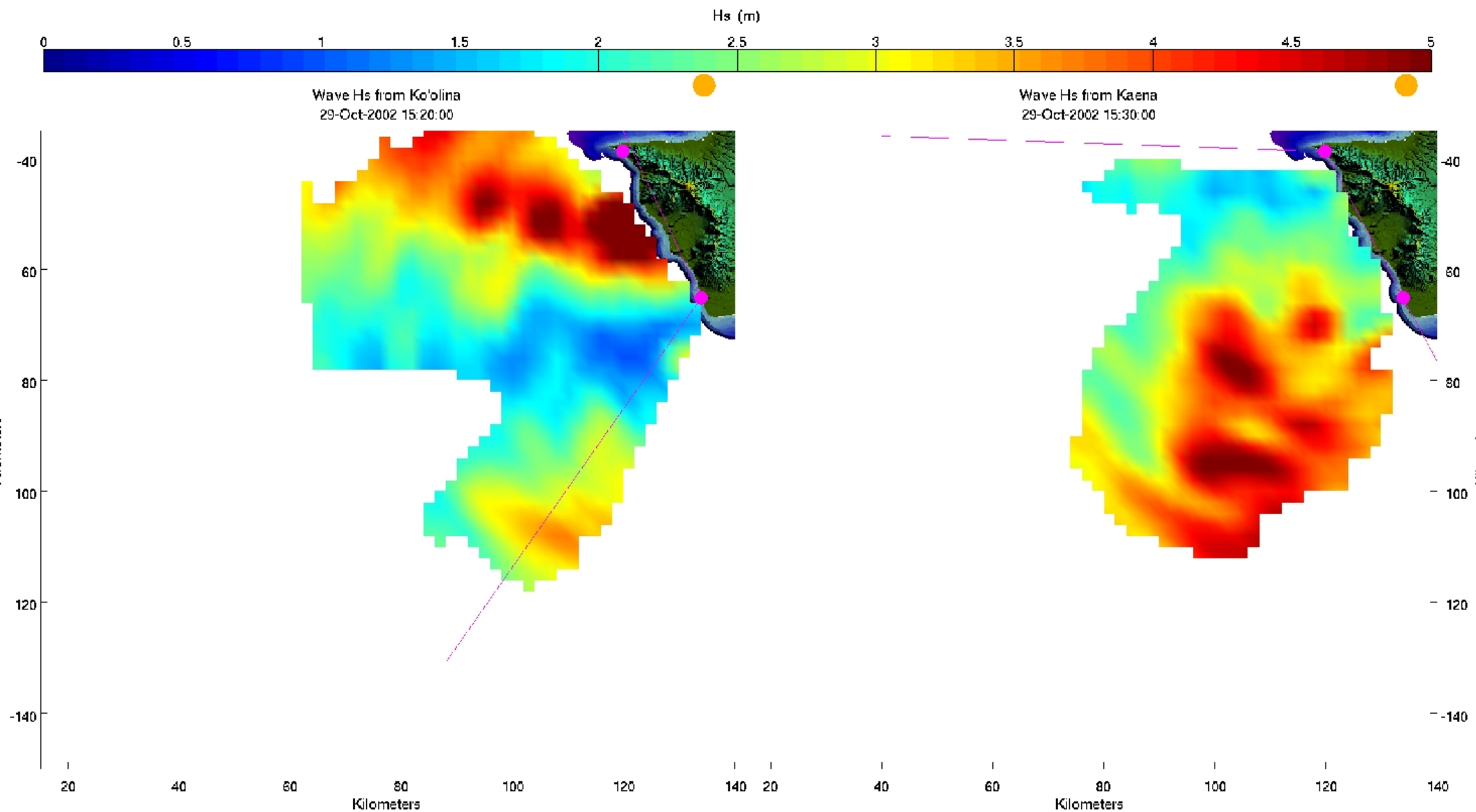
	radar	$\lambda_B$ (m)	Observed $H_{\text{rms}}$ (m)	Correlation	$\sigma$
Maresca and Georges [1980]	WARF*	9-15	0.3-0.8		7,17%
Heron et al. [1985]	COSRAD	5	0-1		> 15%
Heron et al. [1998]	OSCR	6	0-1	0.97	20%
Wyatt and Green [2002]	PISCES, WERA	6-15, 10-25	0.5-4	0.94	10.5%
Wyatt et al. [2003]	WERA		1-8	0.94	16.6%
Wyatt et al. [2005]	OSCR		0.5-3	0.88	
Wyatt et al. [2006]	PISCES	15-21	0.5-8	0.90	33.8%

$\lambda_B$ : Bragg wavelength,  $\sigma$ : normalized standard deviation, i.e. relative standard error

\* skywave radar

**WERA                      9.4                      1- 4                      0.93                      16 %**

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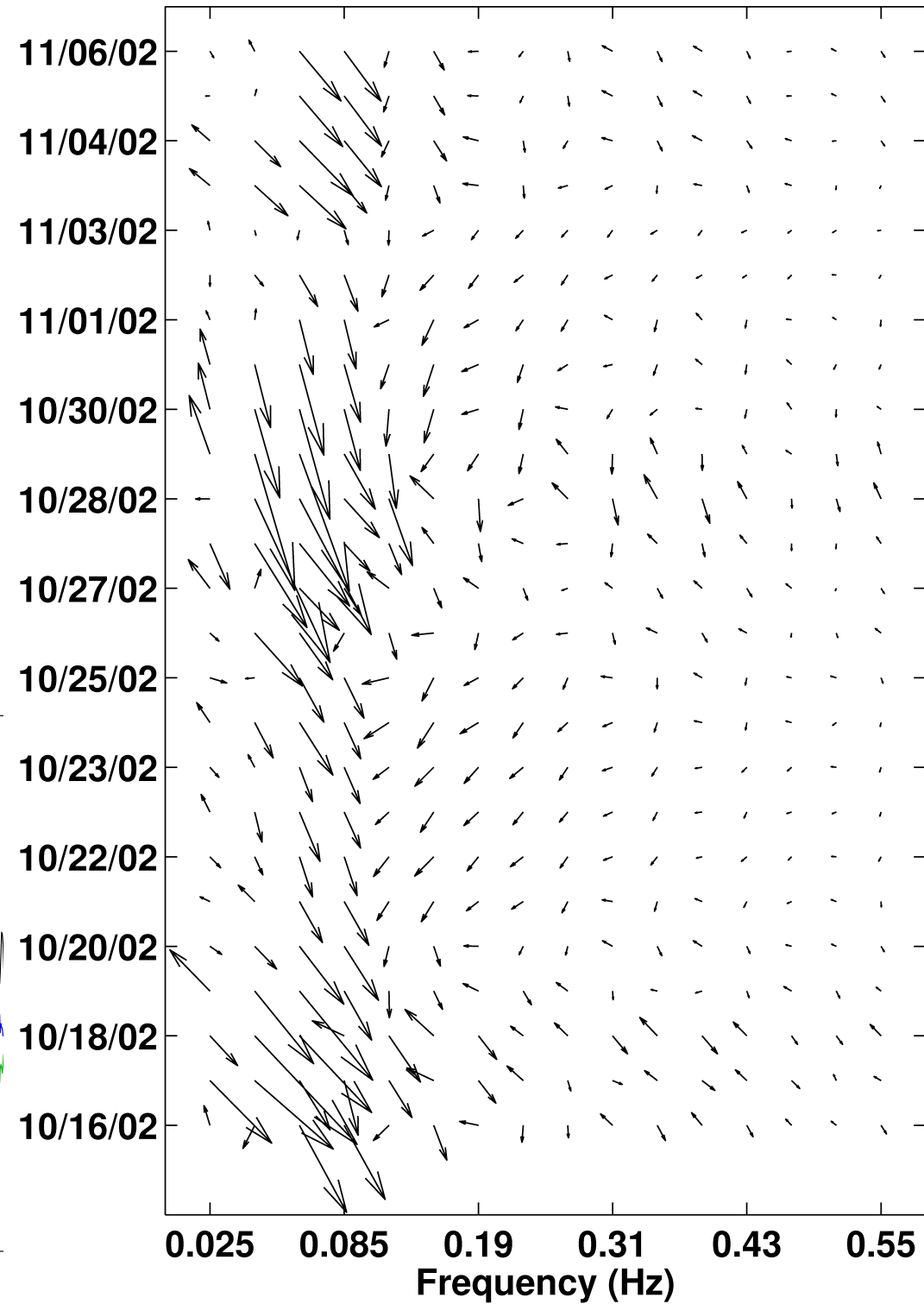
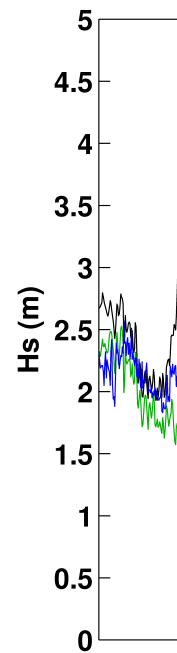
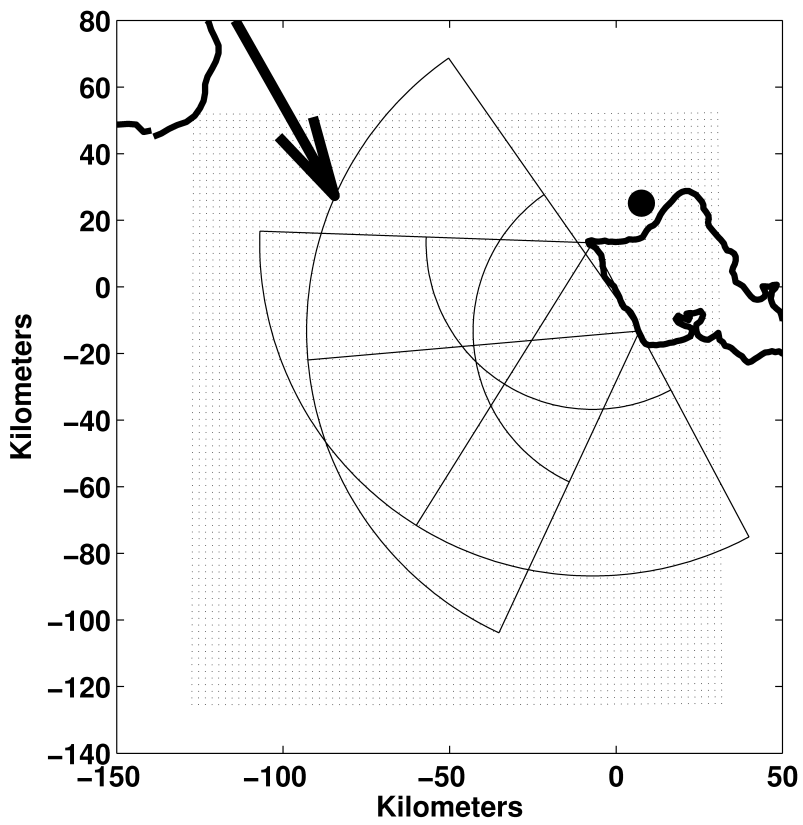
Spatial field of Hs estimates showed large variations in space and time.  $O(5)$  m differences within 25 km and 1 hour

# Outline

1. Theory
  1. First order measurement
  2. Second order measurement
2. Noise
3. CEOF
4. Model-based
5. SNR filtering



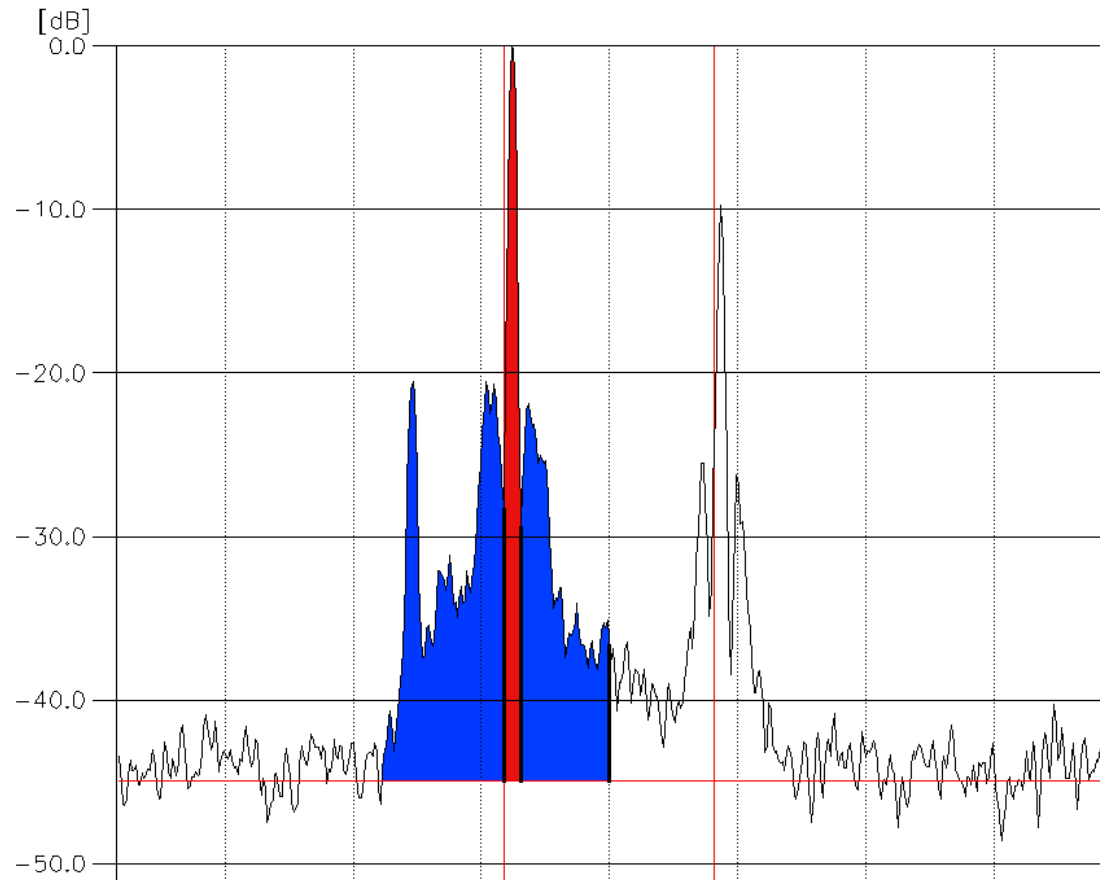
# Land Shadowing



# Significant wave height

$$h_{rms}^2 = \frac{2 \int_{-\infty}^{\infty} \sigma_2(\omega_d) / \Gamma d\omega_d}{k_0^2 \int_{-\infty}^{\infty} \sigma_1(\omega_d) d\omega_d}$$

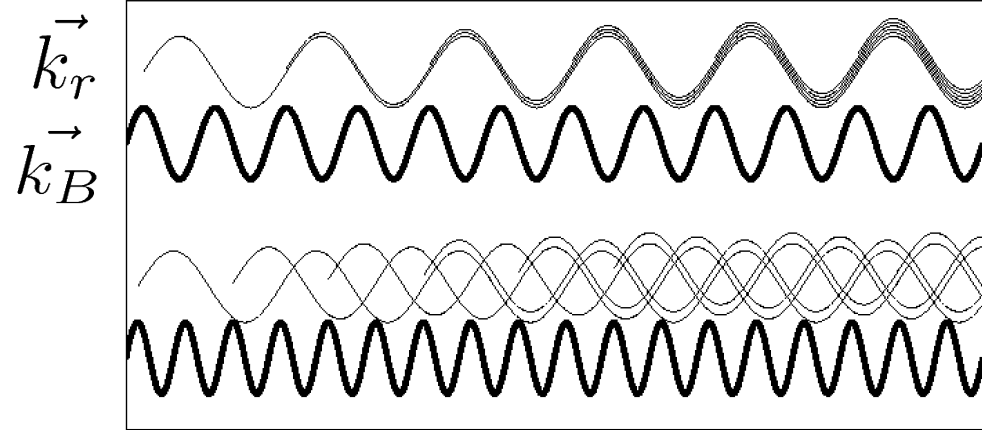
Using theory derived from electromagnetic and hydrodynamic first principles, the radar is capable of measuring  $H_s$  as a ratio of integrated first order (linear) to second order (non-linear) energies.



# First Order Bragg Scattering

To any order, coherent signal will only occur from surface features with the Bragg wavelength

$$\sigma_1(\omega_d) = 2^7 \pi k_0^4 S(\vec{k}_r) \delta(\omega_d - \omega_r)$$

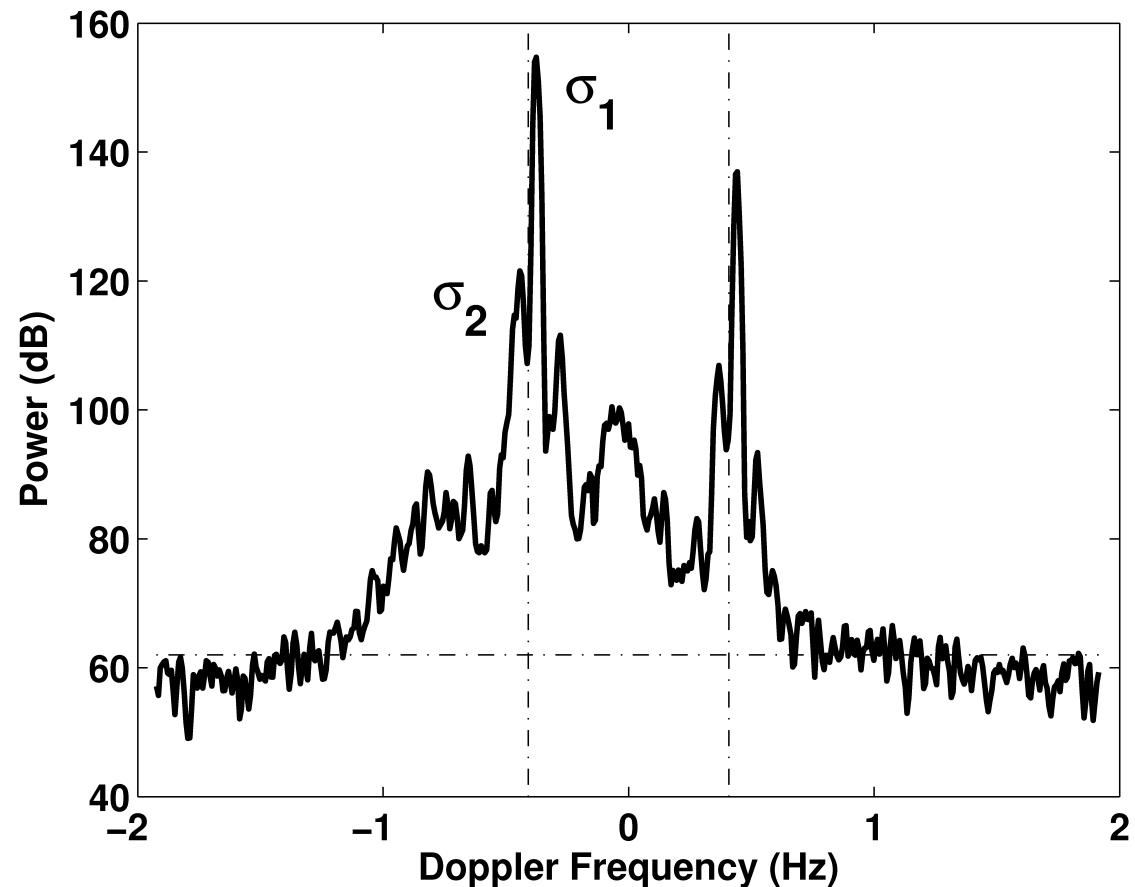


First order scattering is a simple proportional relationship between the EM and ocean waves.

Only *ONE* ocean wavelength causes coherent scattering. It is a known constant.

$$2\vec{k}_r = \vec{k}_B$$

$$\sigma_1(\omega_d) \quad \omega_d$$





# Second Order Scattering

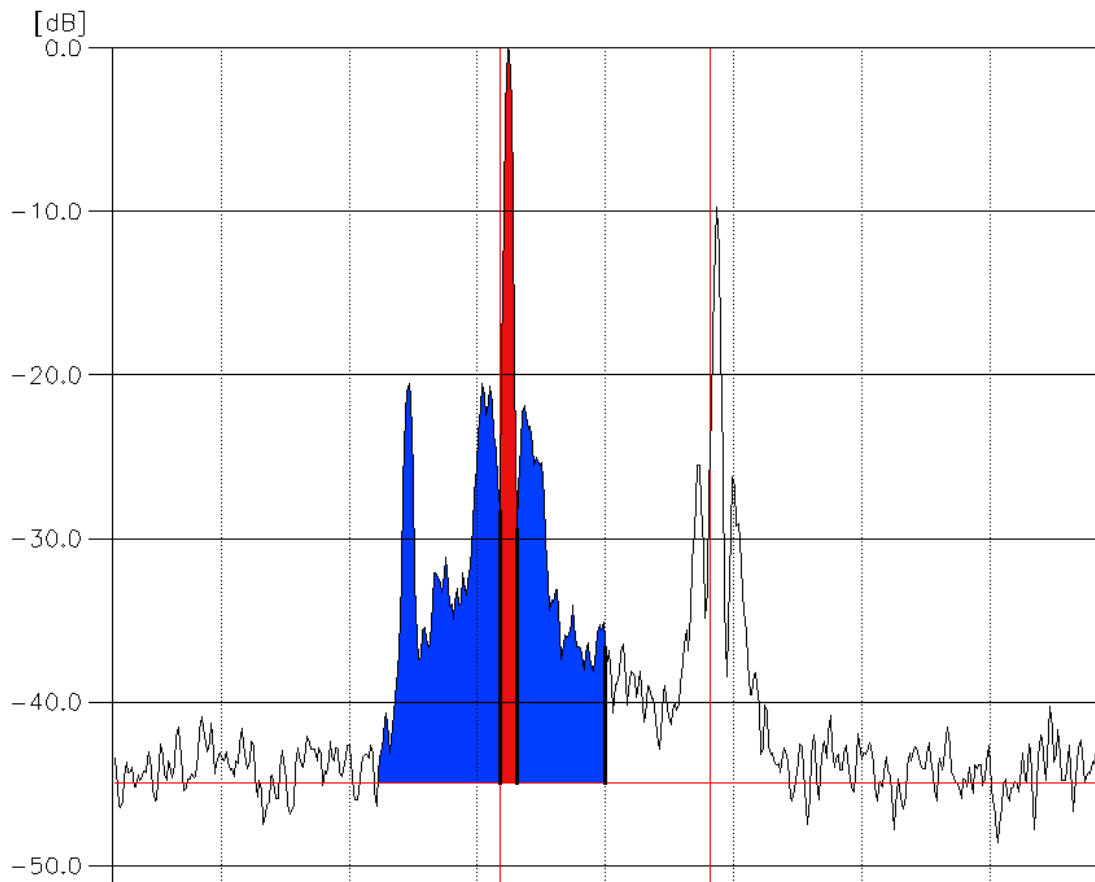
$$\sigma_2(\omega_d) = 2^4 \pi k_0^4 \int \int_{-\infty}^{\infty} |\Gamma(\vec{k}_1, \vec{k}_2)|^2 S(\vec{k}_1) S(\vec{k}_2) \delta(\omega_d - \omega_1 - \omega_2) d\vec{k}$$

$$\vec{k}_B = \vec{k}_1 + \vec{k}_2$$

$$\omega_d = \omega_1 + \omega_2$$

Information encoded  
primarily in  
amplitude

Again, Second-Order signal  
can only occur for coherent  
scattering off features with  
the Bragg wavelength.



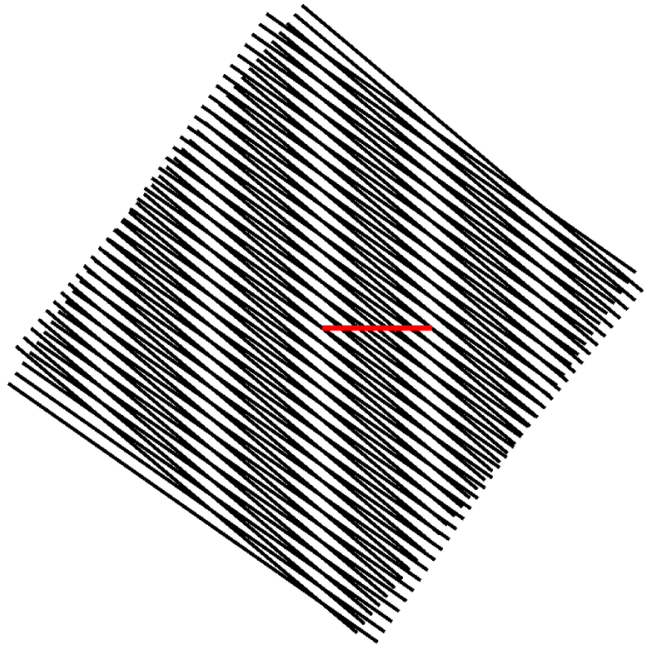
These are *nonlinear ocean waves*, and EM double-scattering

Perturbation expansion of nonlinear equations for free water surface and the EM fields yields second-order terms.

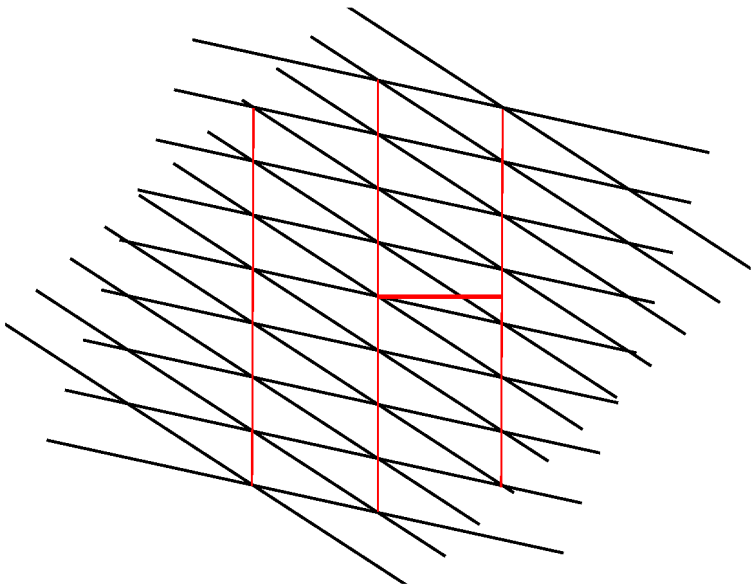
Represented as coupling coefficient Gamma.

$$\Gamma = \Gamma_H + \Gamma_E M$$

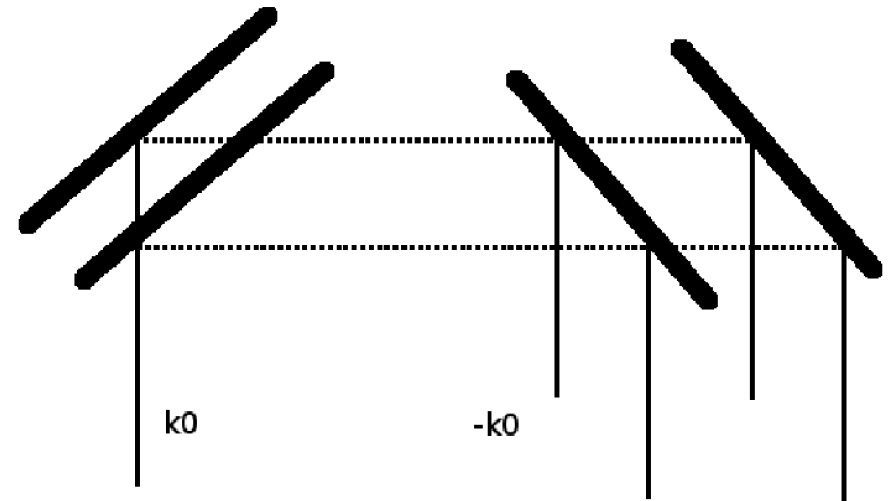
Coupling Coefficient, i.e.  
Second-Order Weighting  
Function



$$\vec{k}_B = \vec{k}_1 + \vec{k}_2$$



Hydrodynamic



Electromagnetic

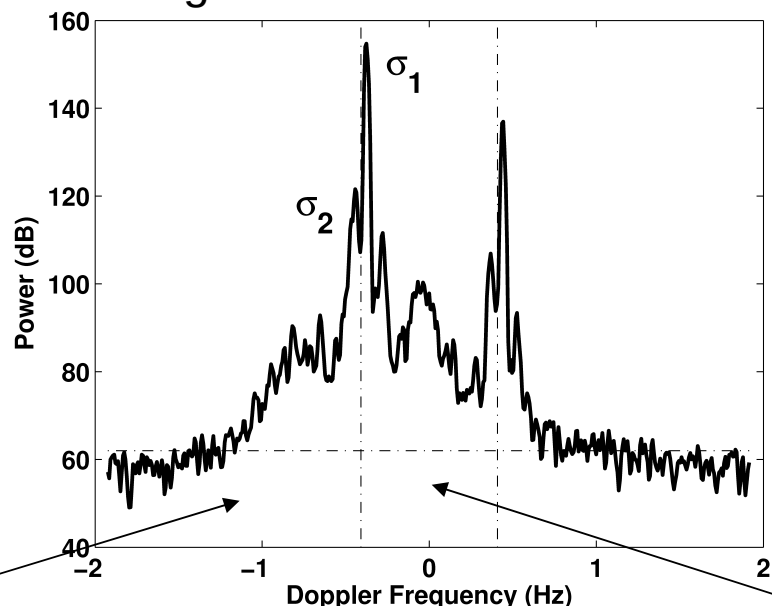
$$\sigma_2(\omega_d) = 2^4 \pi k_0^4 \int \int_{-\infty}^{\infty} |\Gamma(\vec{k}_1, \vec{k}_2)|^2 S(\vec{k}_1) S(\vec{k}_2) \delta(\omega_d - \omega_1 - \omega_2) d\vec{k}$$

The second order integral relationship is a convoluted mapping of the ocean wave directional spectrum to the radar spectrum.

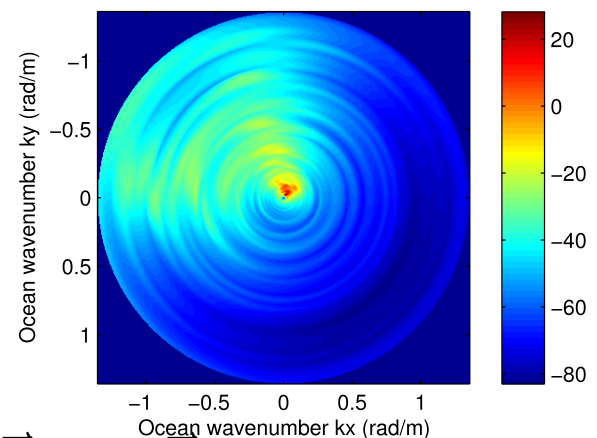
Each radar frequency is an integration over a closed wavevector contour.

$$\omega_d = \omega_1 + \omega_2$$

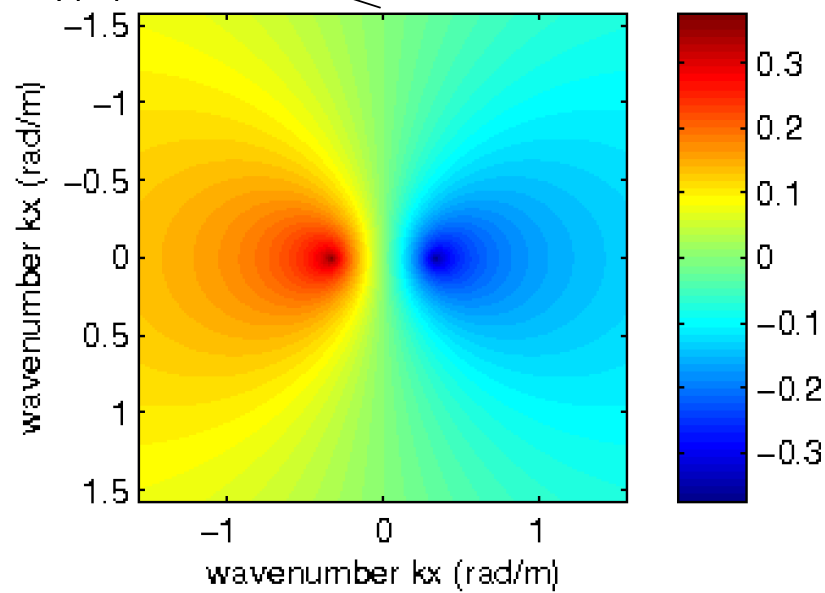
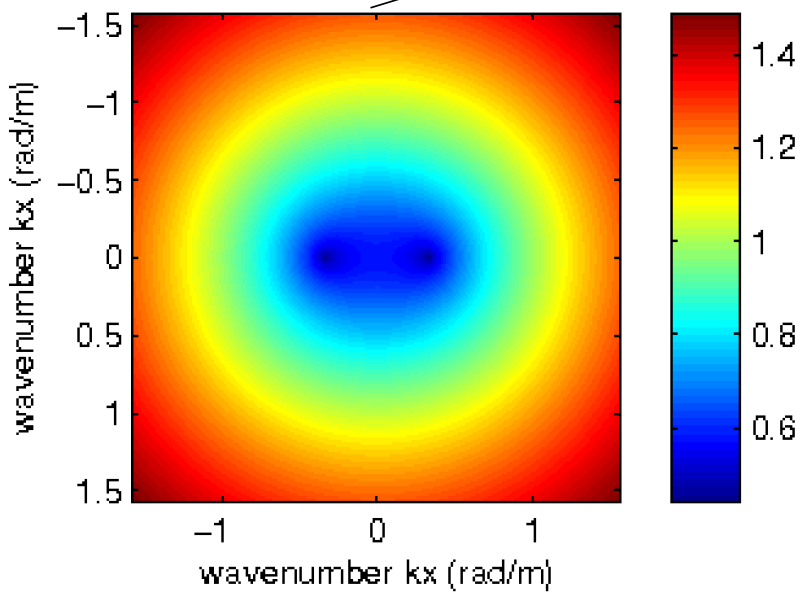
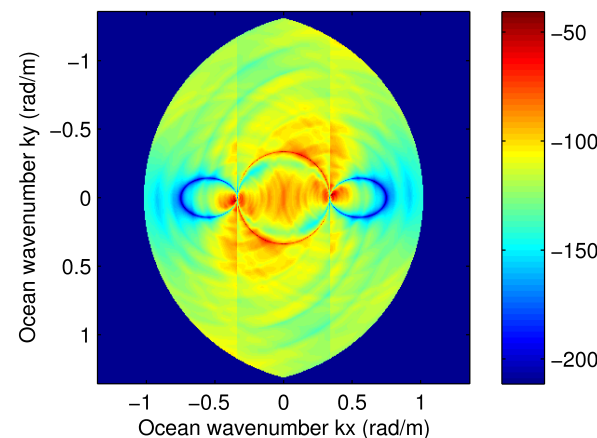
$$\vec{k}_B = \vec{k}_1 + \vec{k}_2$$



$$S(\vec{k})$$

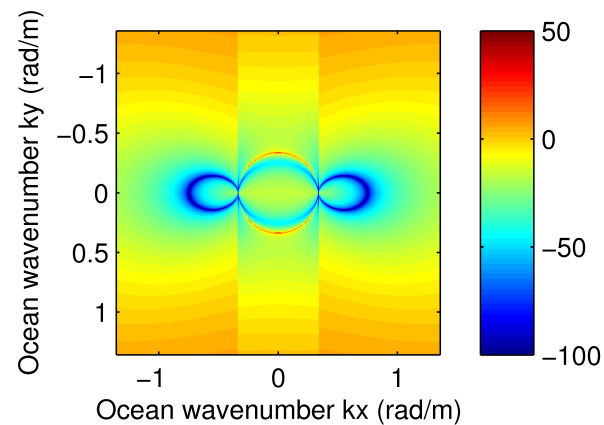
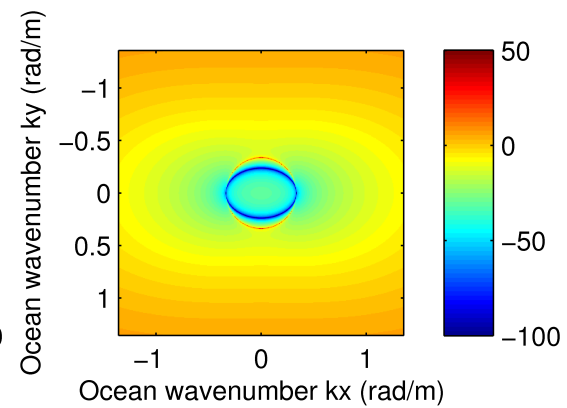
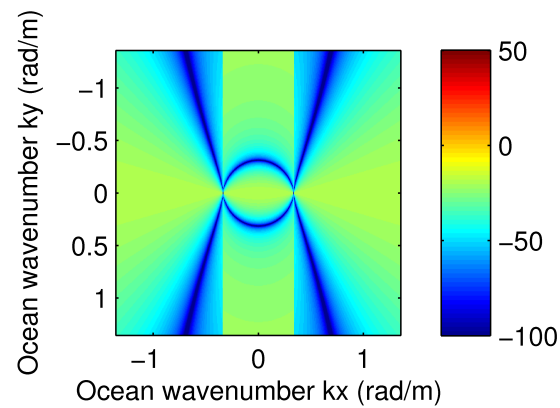
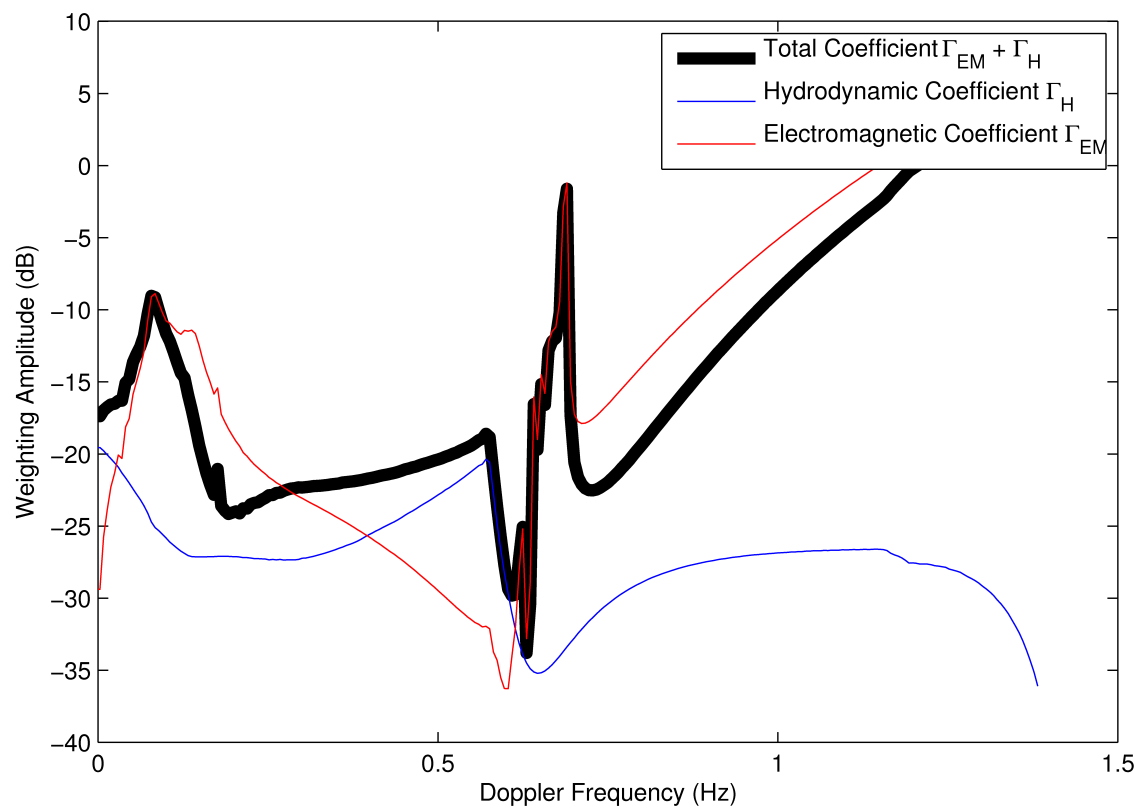


$$|\Gamma|^2 S(\vec{k}_1) S(\vec{k}_2)$$





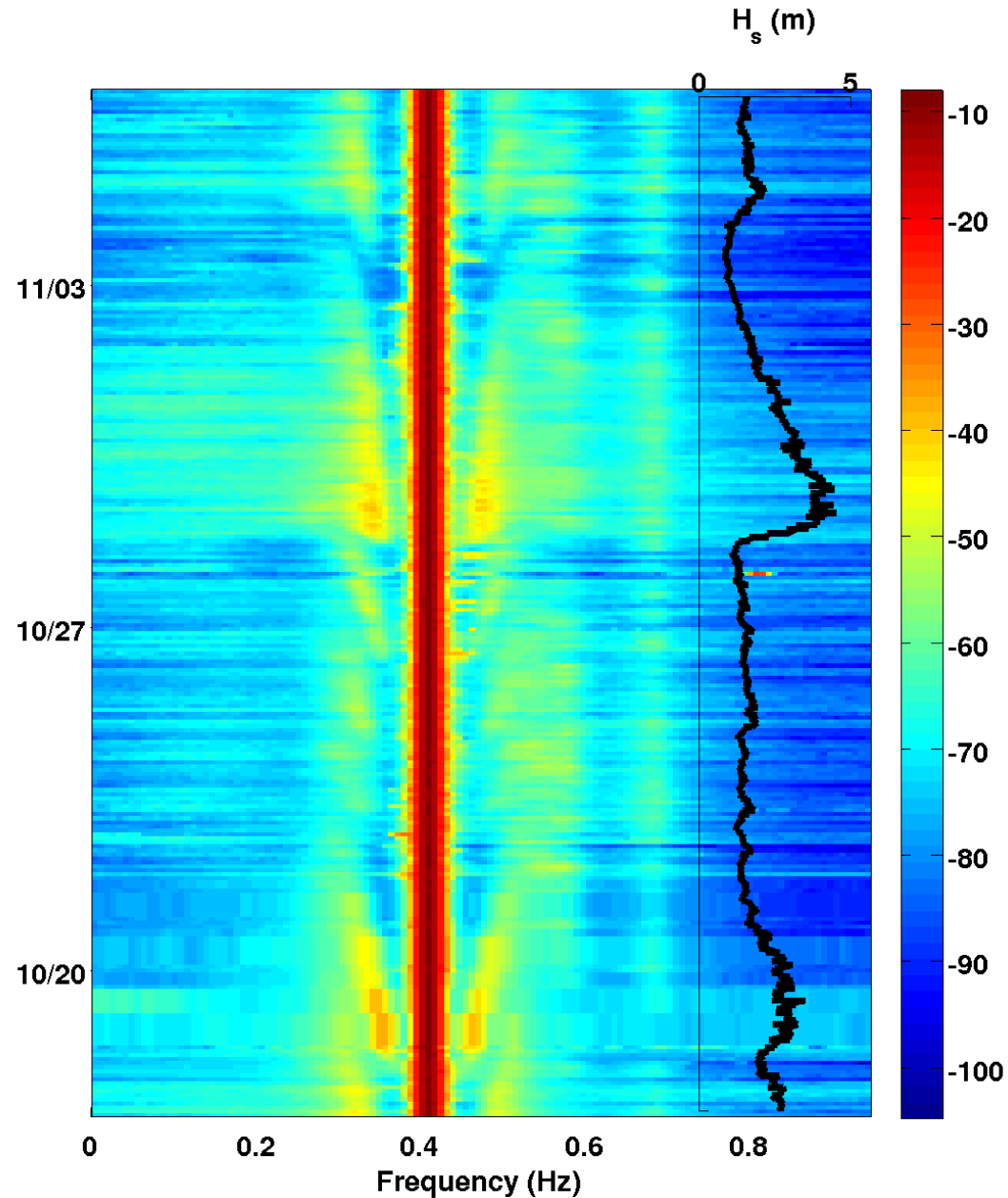
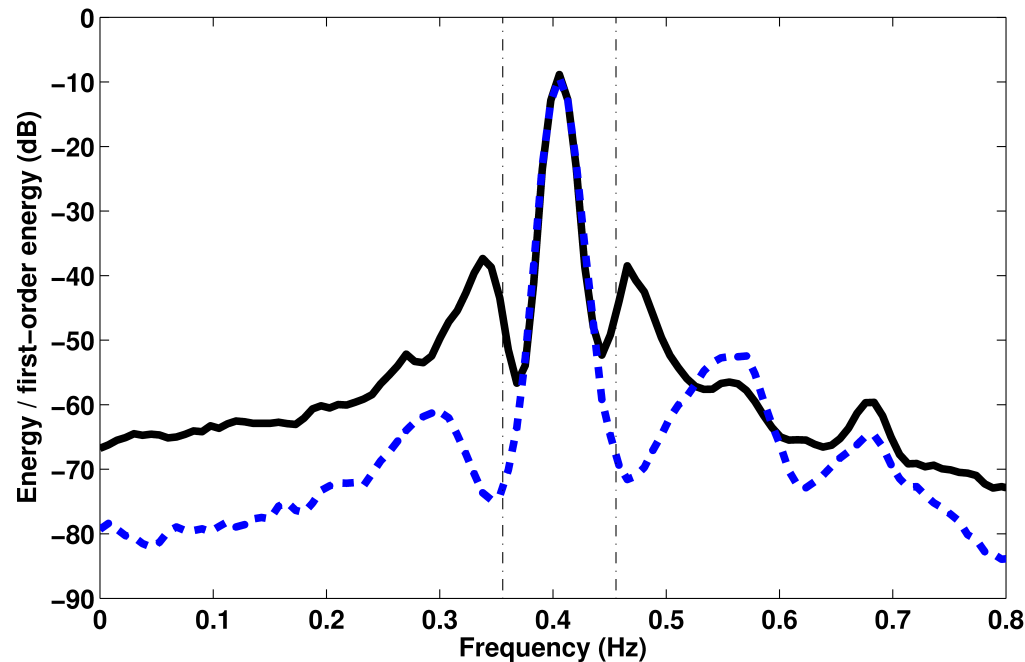
$$\Gamma = \Gamma_H + \Gamma_{EM}$$

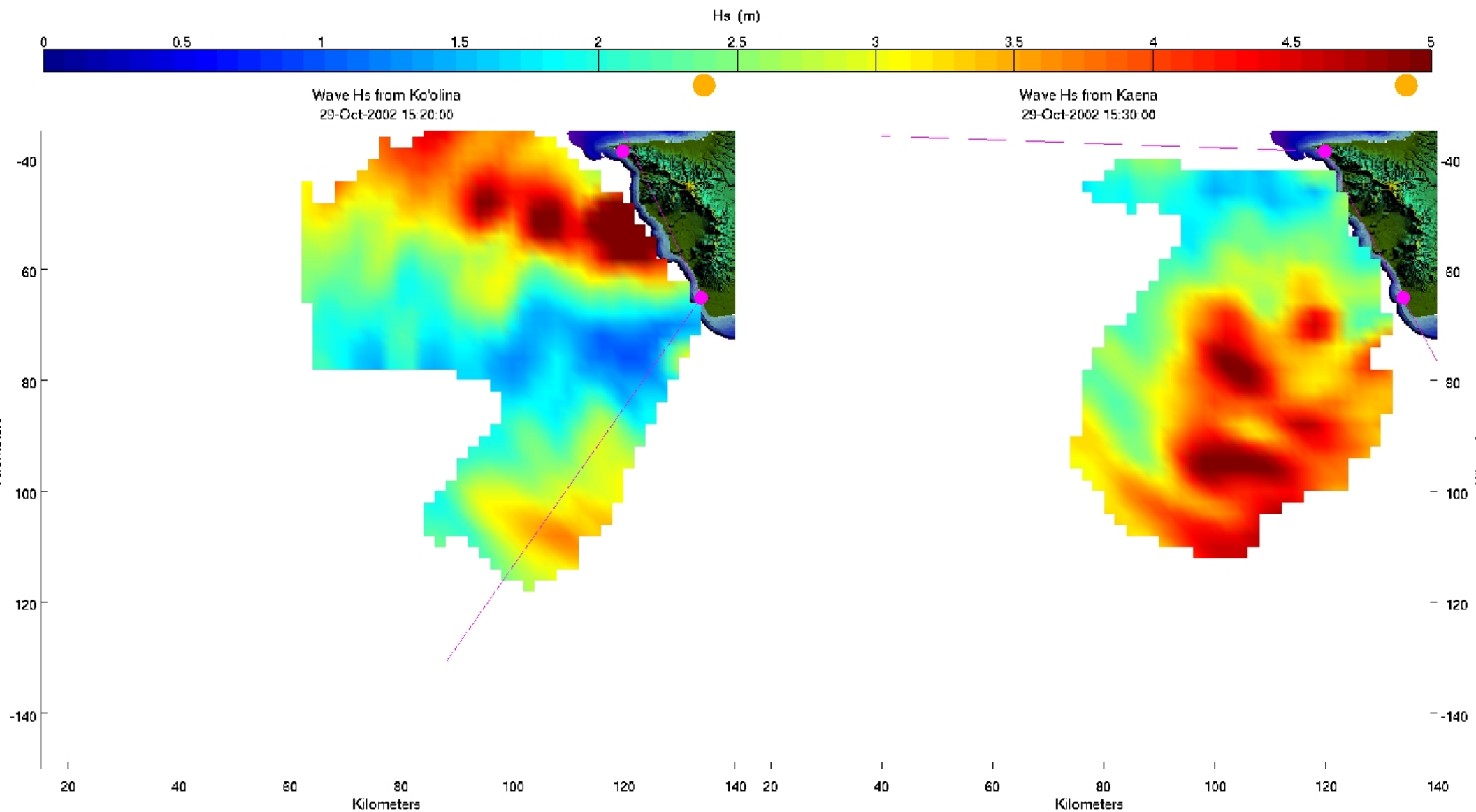


# Waveheight bias

Scattering theory predicts the frequency separation between first and second order peaks will decrease as the ocean swell moves to lower frequencies.

As the energy of the peaks is convolved, their ratio is biased towards one, i.e. the  $H_s$  is biased low.

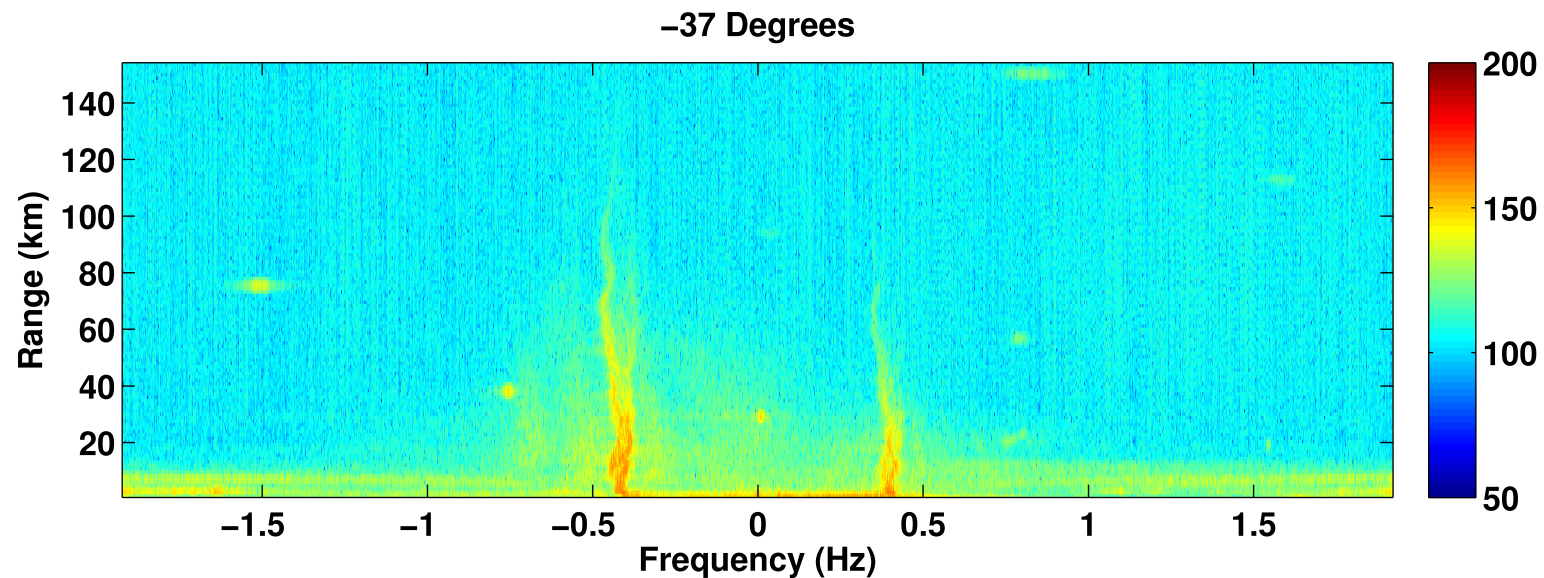
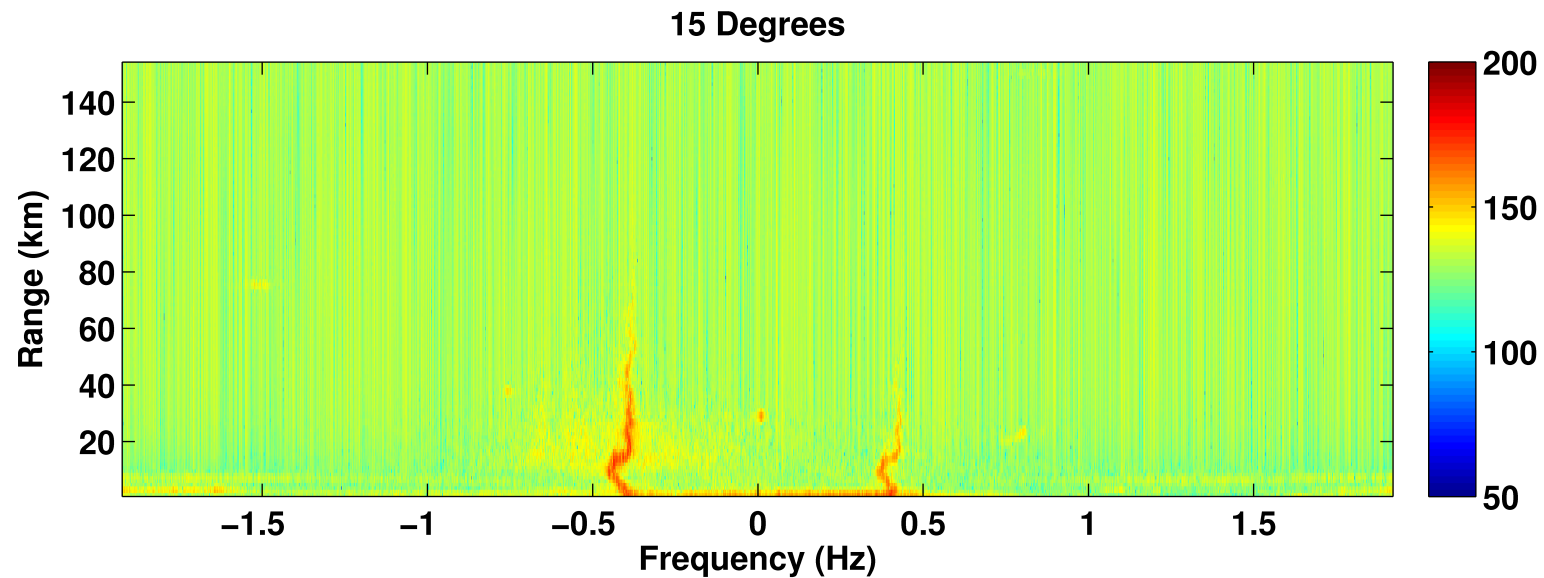




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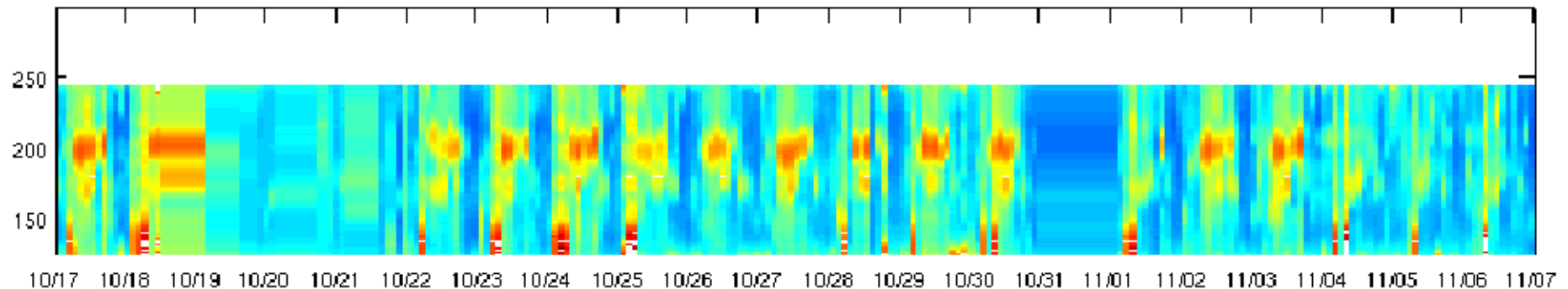
- Noise is contaminating spectra
- Range-independence is characteristic of external interference, not systematic



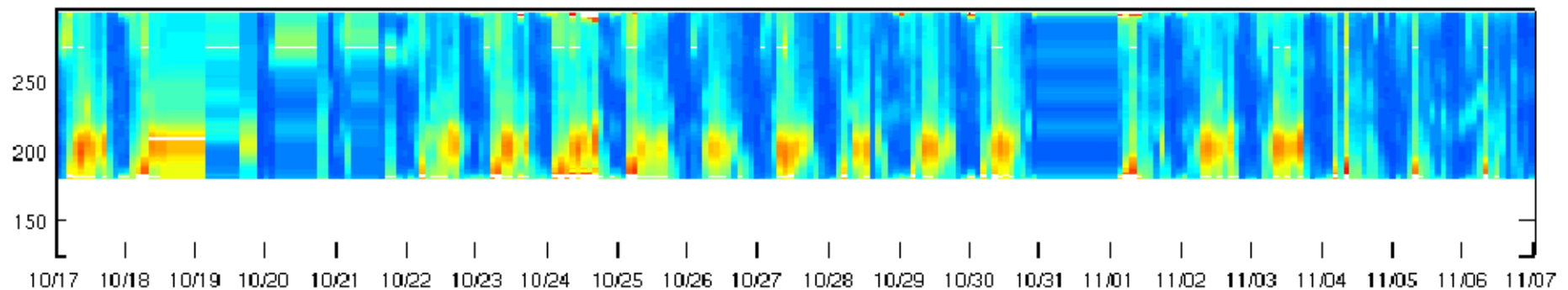
# Noise level timeseries as a function of angle

- Both sites have strong interference at 200 deg incident

Koolina, Angular fcn

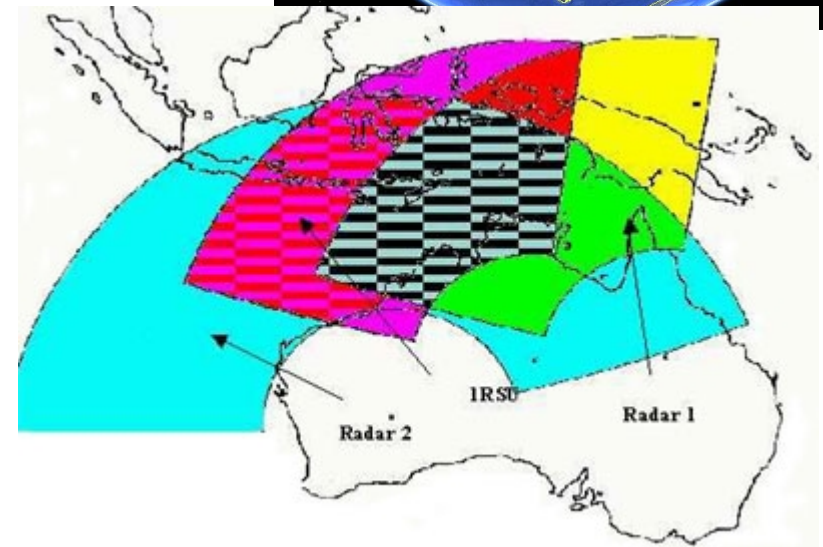
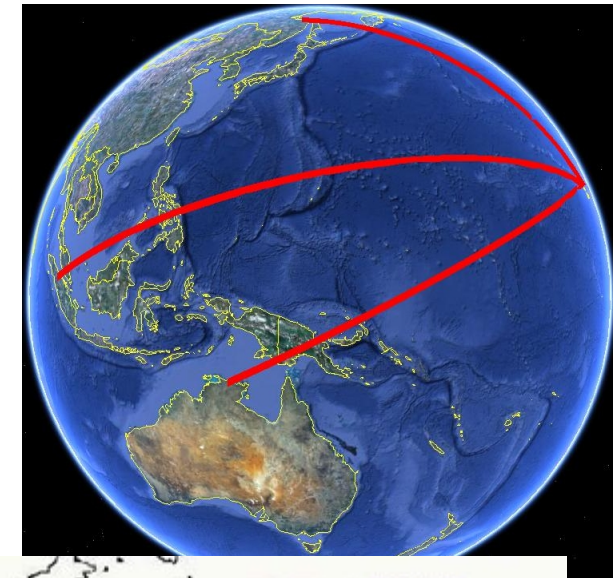
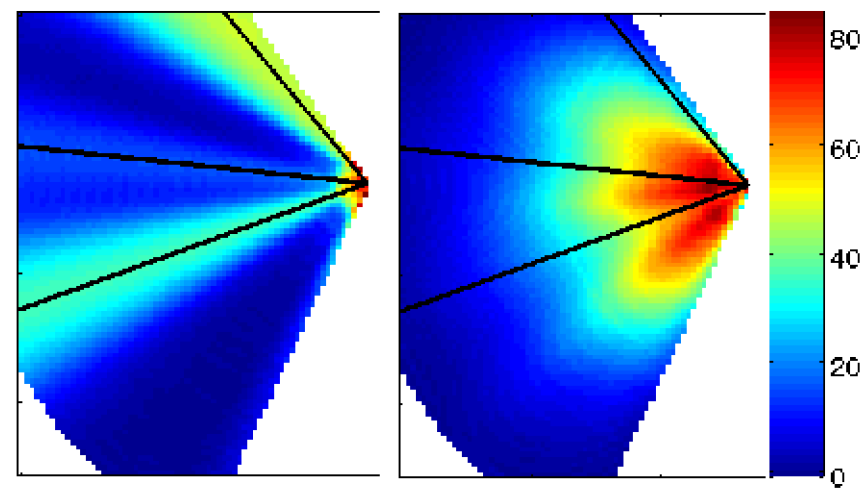


Kaena, Angular fcn

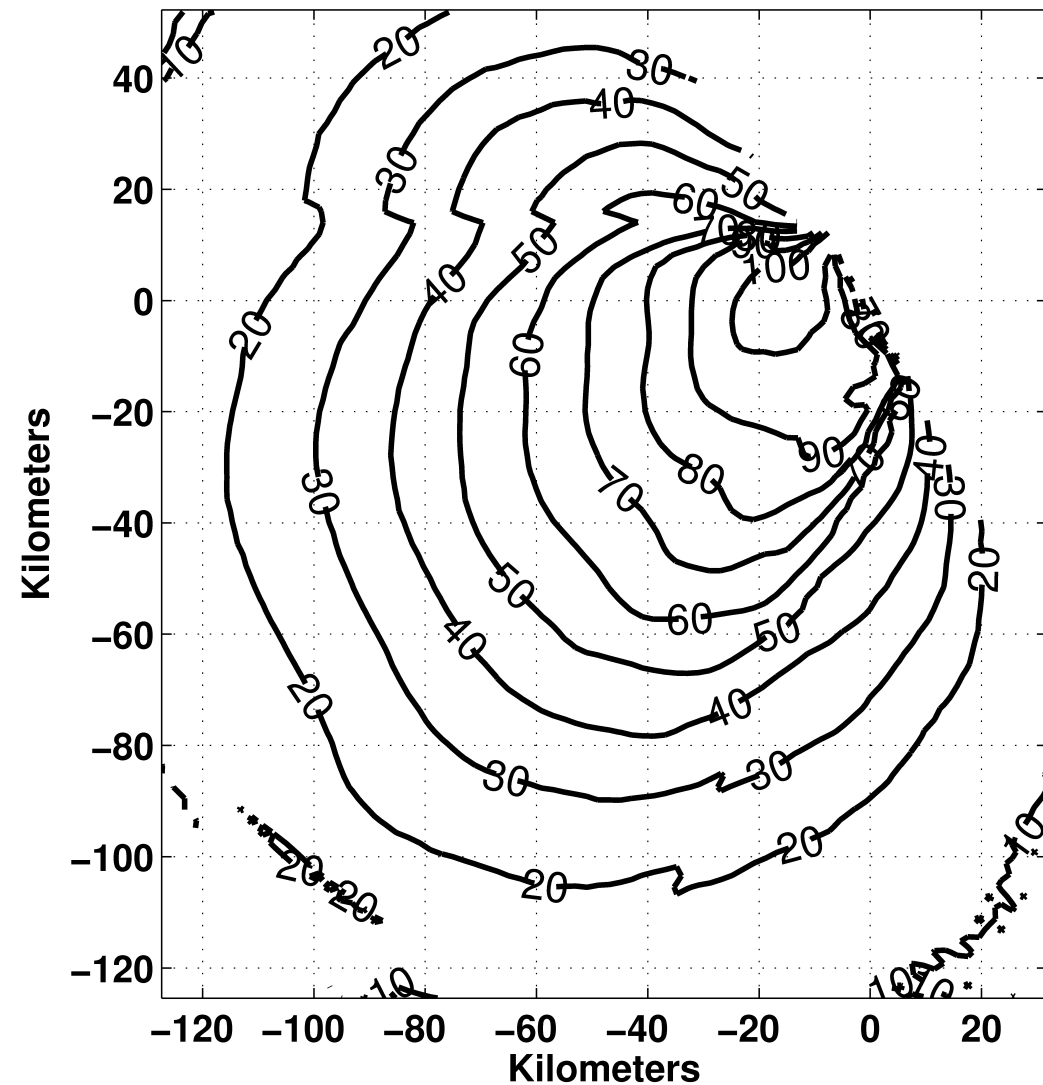
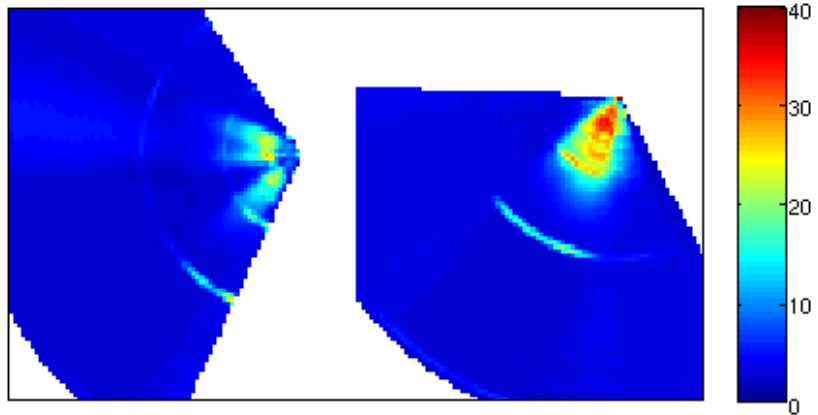
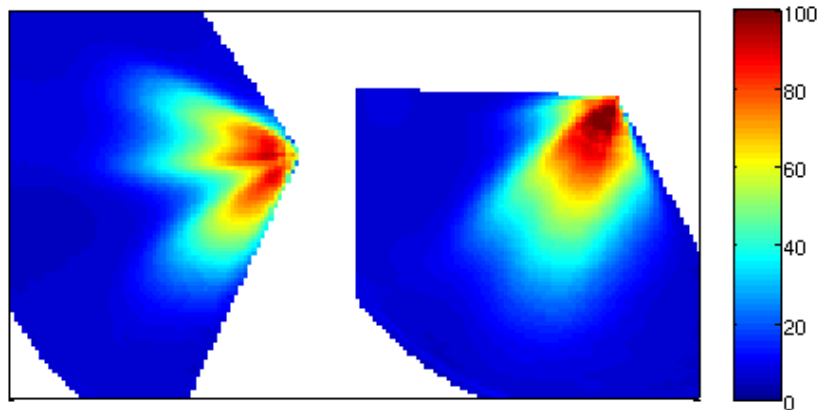


# External Interference

- 
- 
- 
- 
- 



# Signal-To-Noise Ratio (SNR)

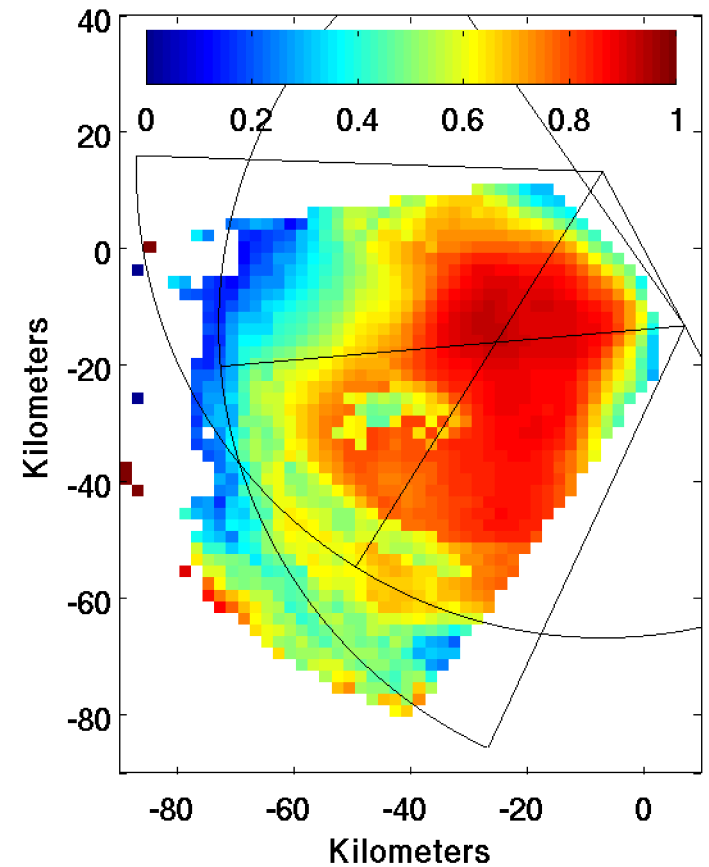
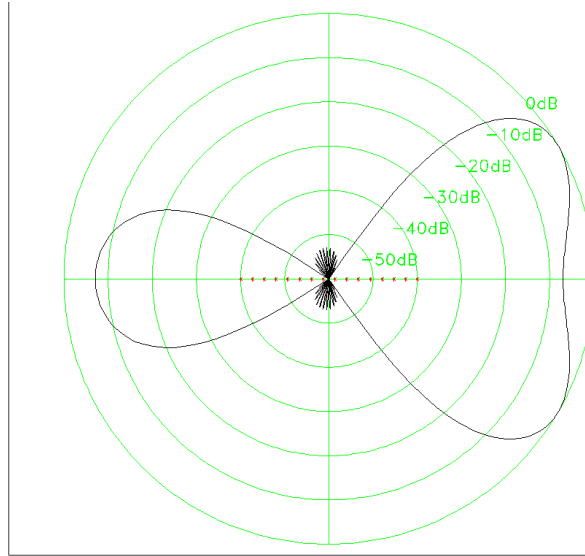
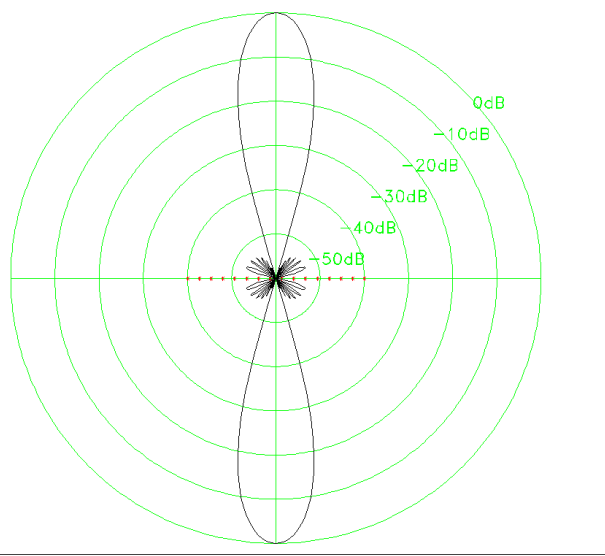
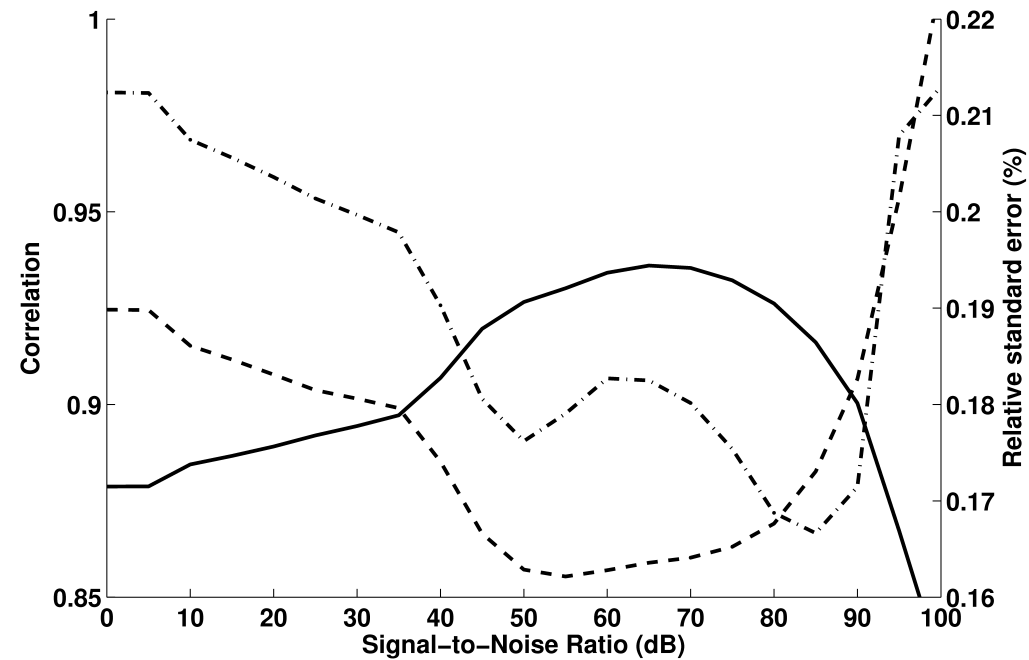


- Mean SNR of 50 dB to ~ 70 km
- Second order energy mostly degraded

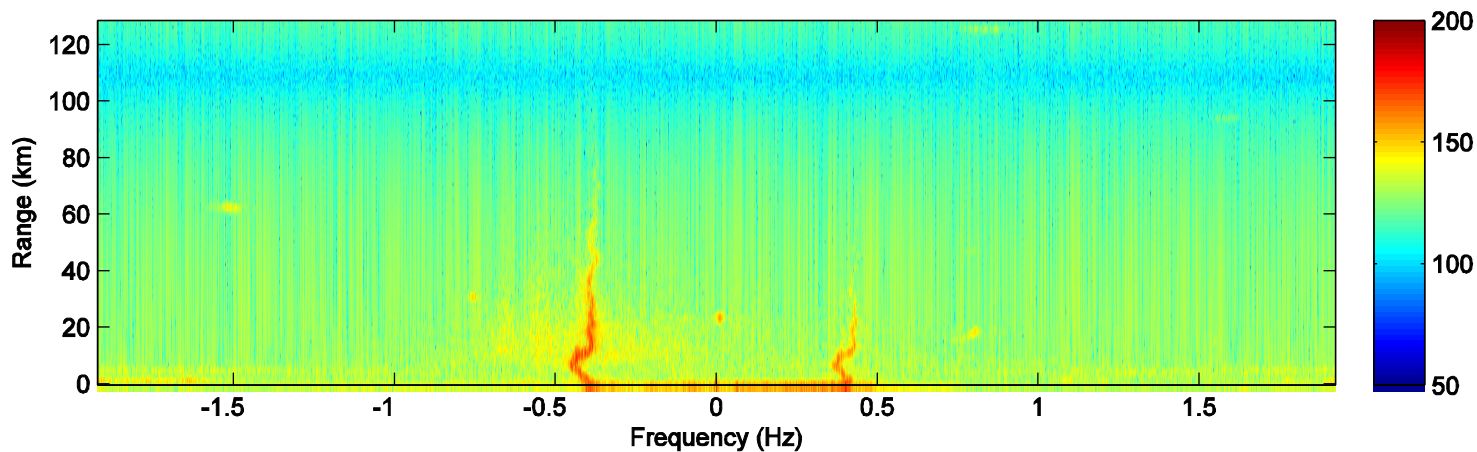
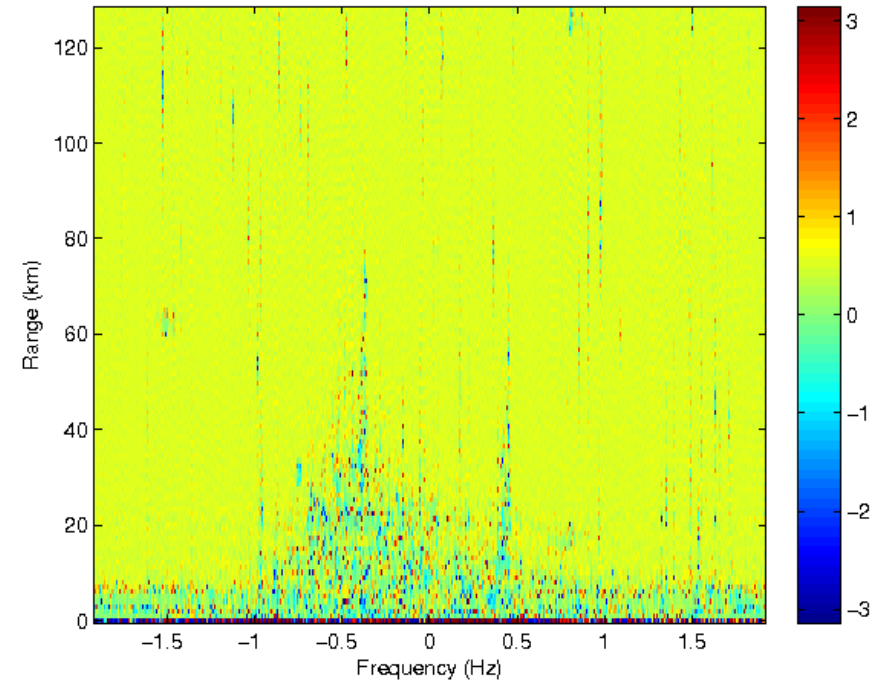
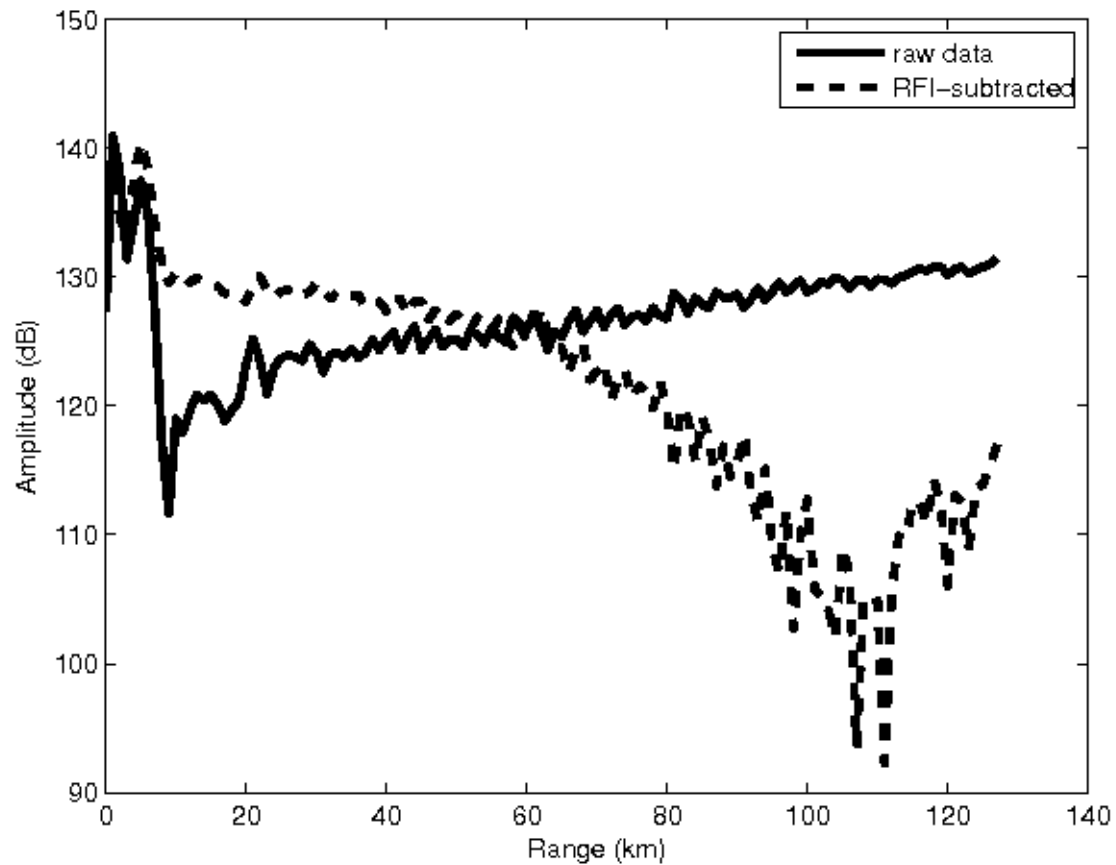


# SNR Filtering

- Small improvement to accuracy:
  - + 0.05 r, - 3% normalized error
- Maximum accuracy occurred at intersection of center-beams; attributed to optimal beamforming
- 80% reduction is usable data
- *Spatial accuracy is not significantly improved*



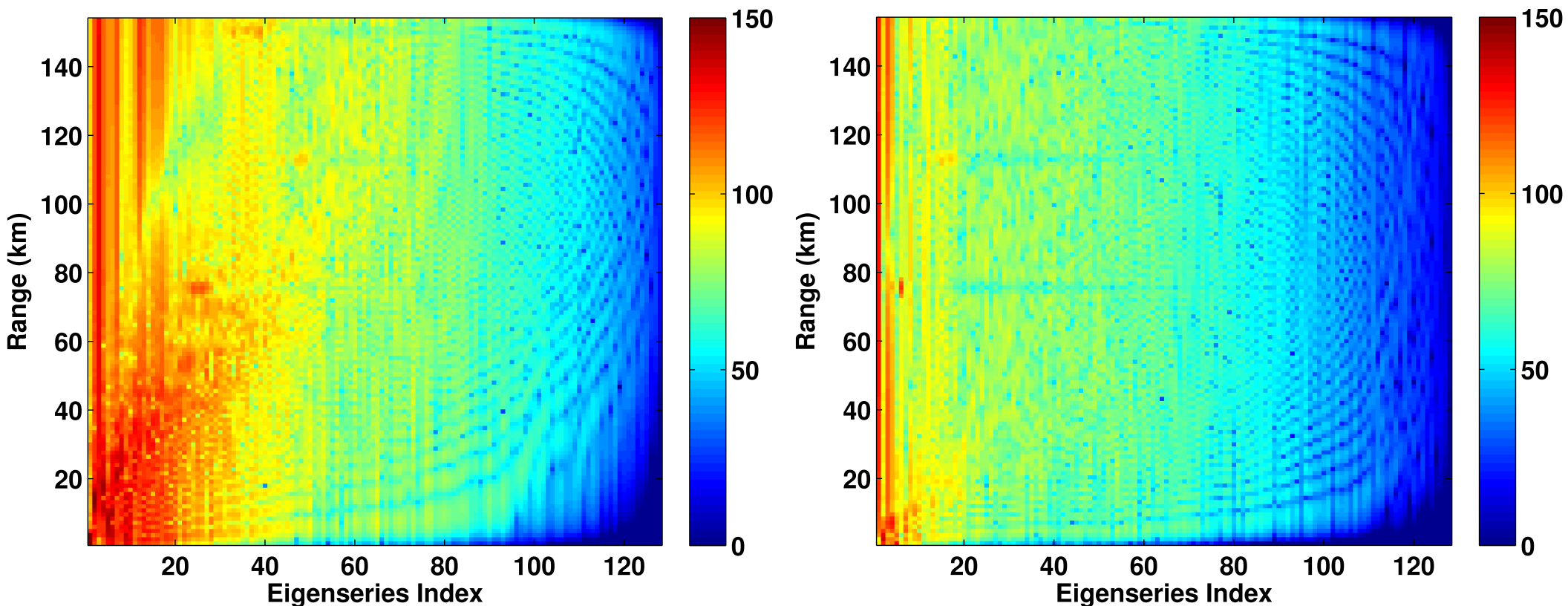
# Model-based RFI reduction



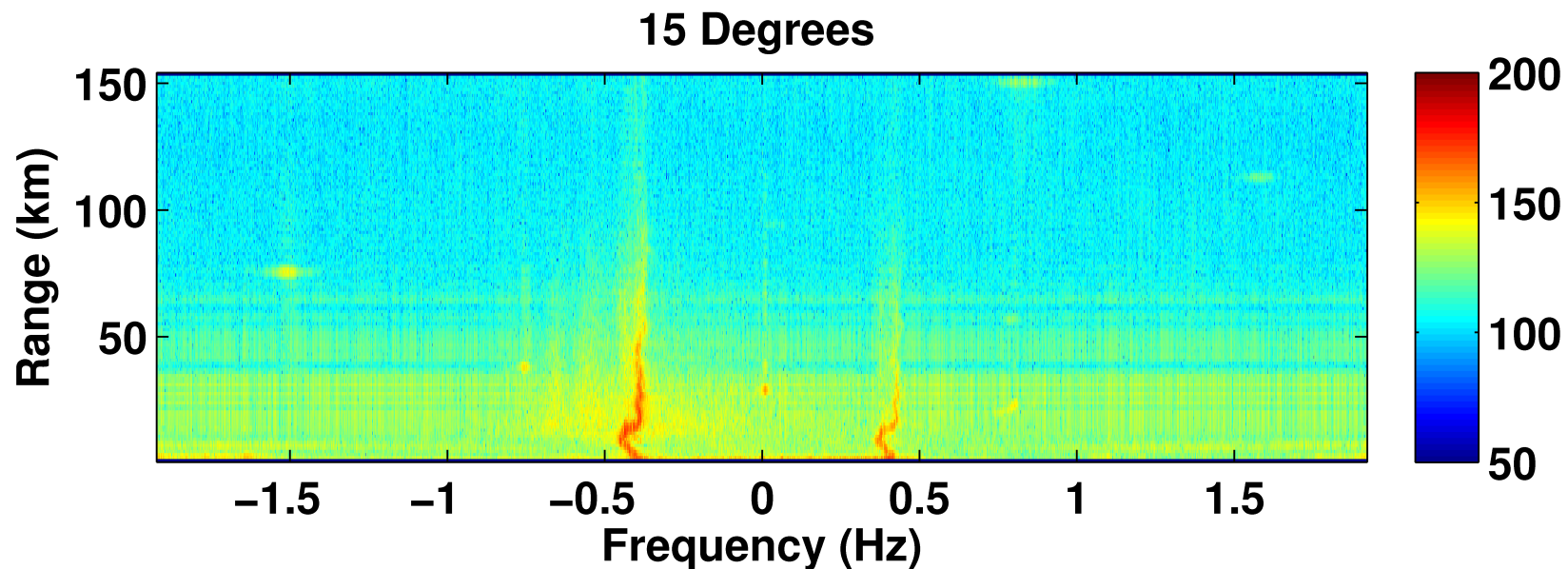
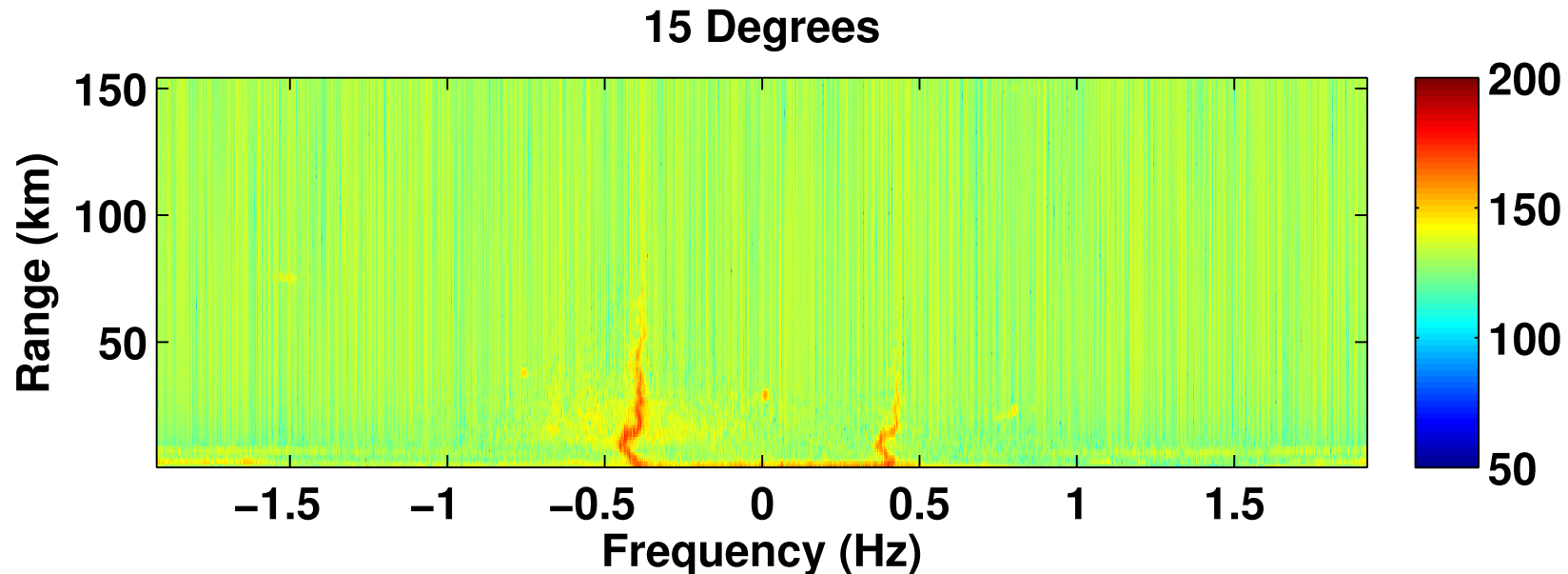
# Complex Eigenfunction Filtering

Eigen-decomposition methods are often based on the assumption that signal (or noise) can be represented by only a few independent modes

Energetic noise and signal modes are not independent.



Eigenfunction filtering is only accurate when the signal and noise modes are independent



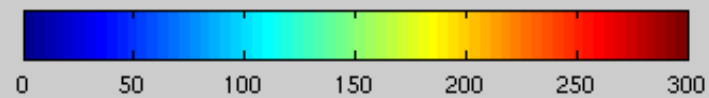
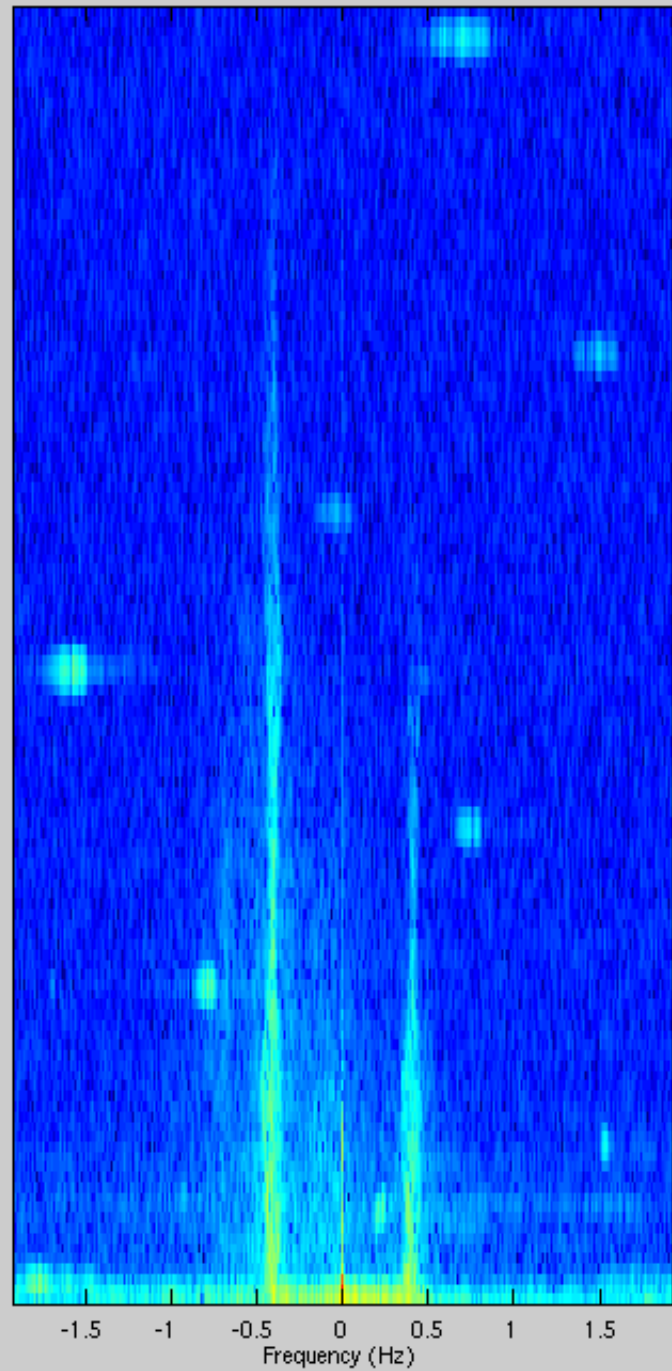
# Acknowledgments

- Mark Merrifield
- Pierre Flament
- Doug Luther
- Cedric Chavanne
- Francois Ascani
- Jerome Aucan
- Derek Young

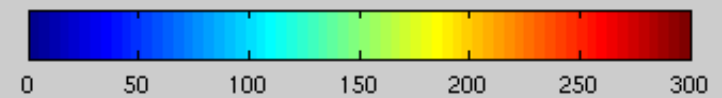
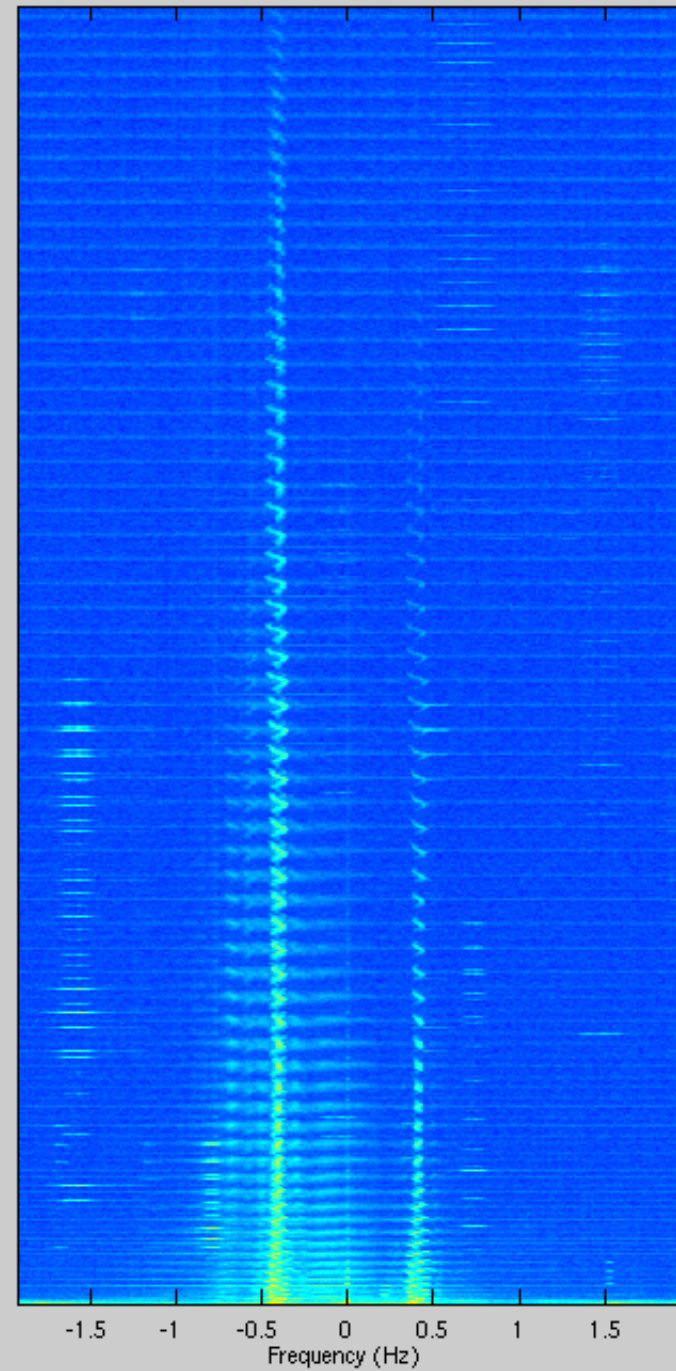




dB PSD  
28-Oct-2002 15:20:00



dB PSD  
28-Oct-2002 15:20:00



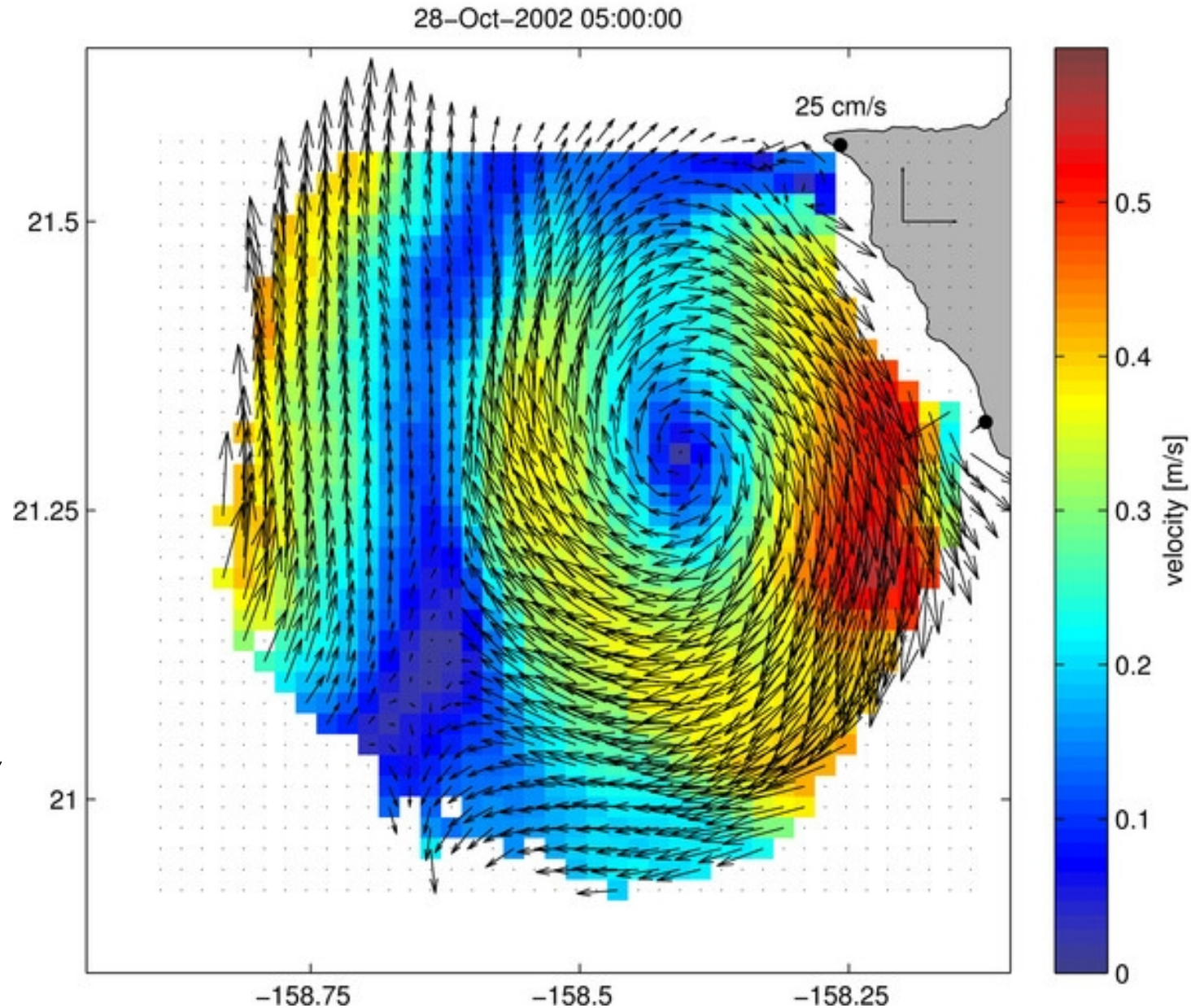
# Surface currents derived from First Order scatter

Chavanne, 2007

Current information is carried in the frequency content of the signal, not amplitude

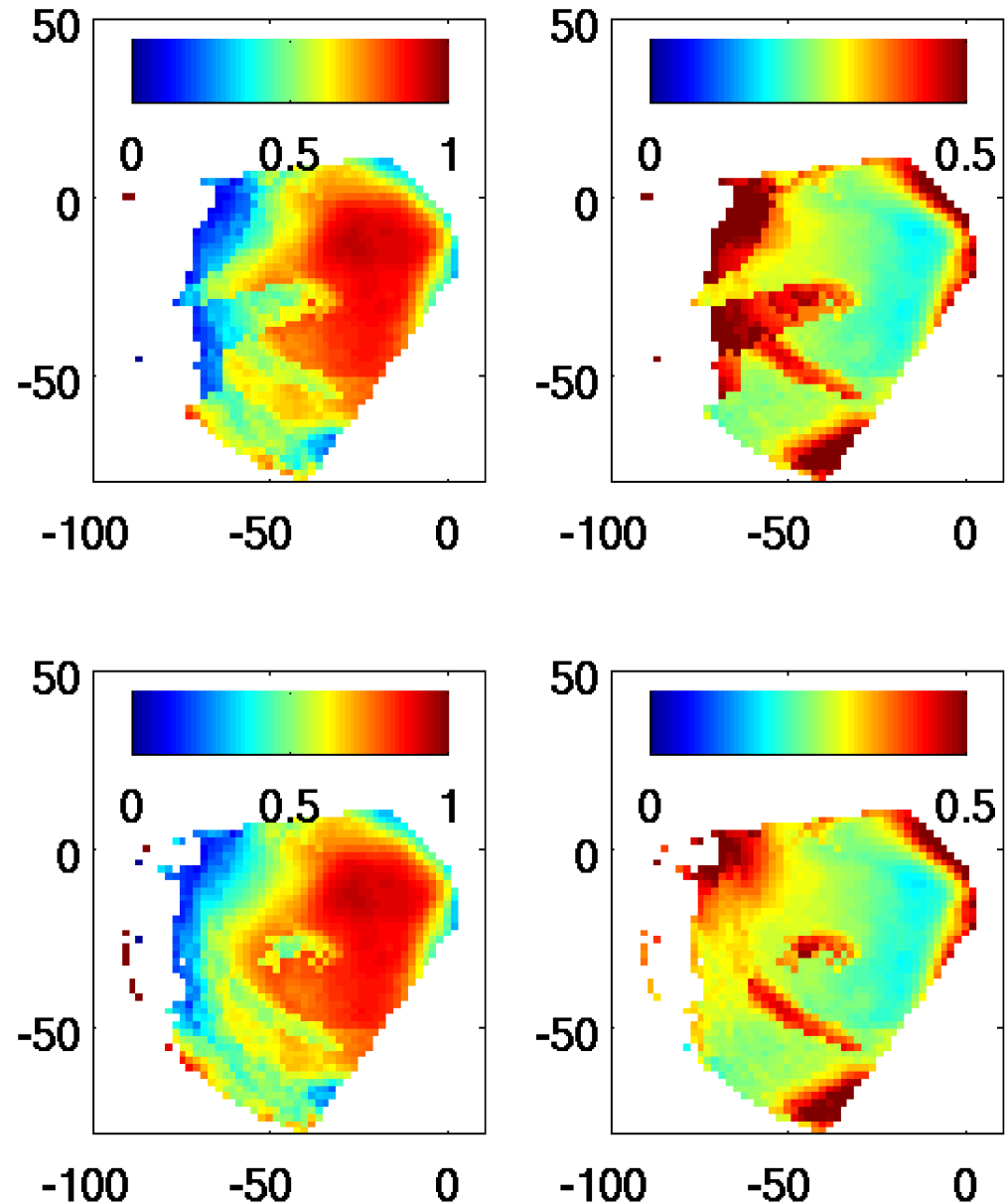
Difference from known frequency is due to surface currents (and Stokes Drift)

$$\Delta c = \sqrt{\frac{g\lambda}{2\pi}} - \lambda f_m$$

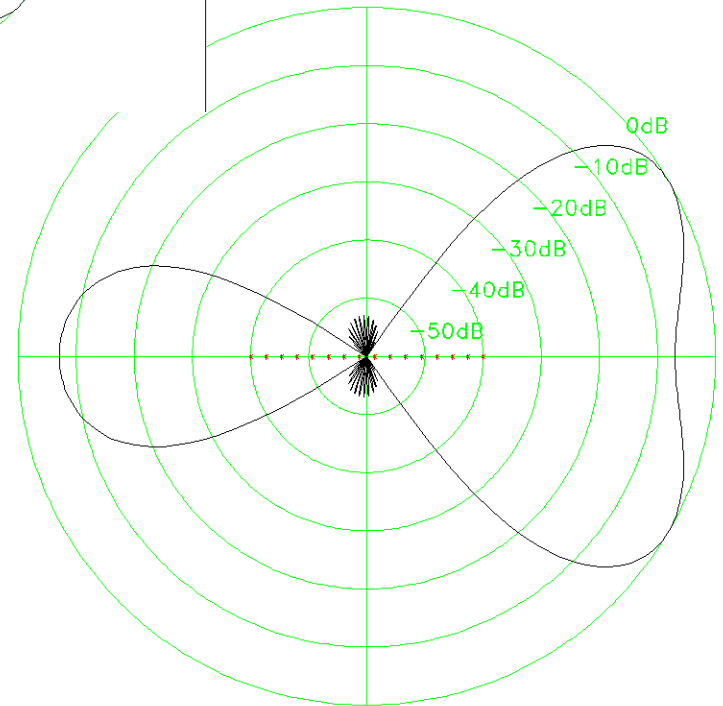
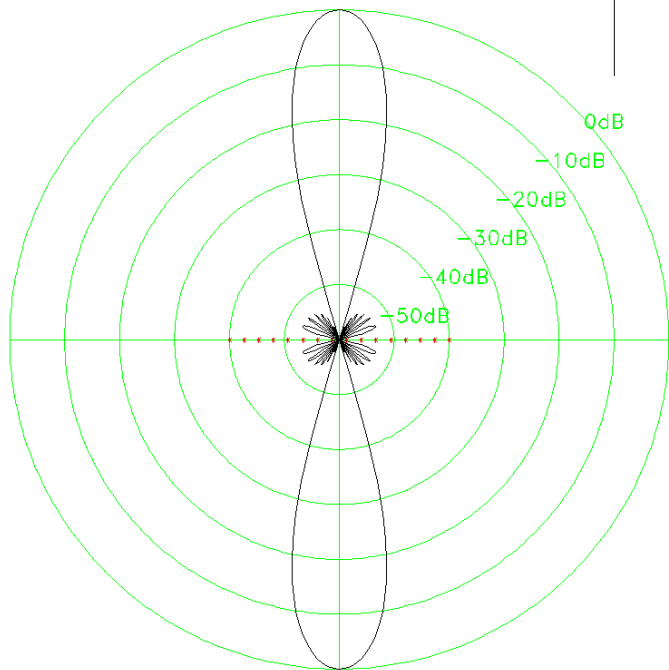
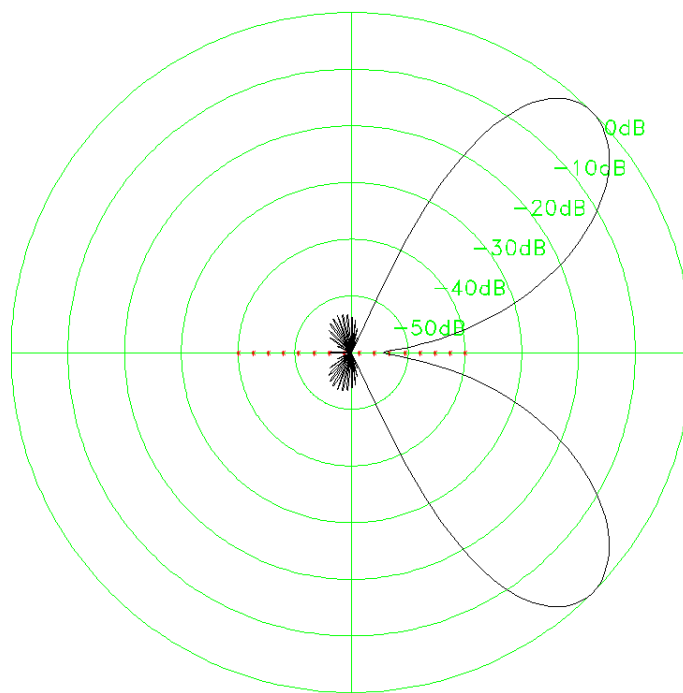


# SNR Filtering: before and after

- Spatial accuracy is not significantly improved



# Beamforming



# Data Products

- Currents
- Significant wave height
- Wind direction
- Wave direction

# Physical Processes

- **Current** measurements cover broad band of motions:  
  
subinertial (<48 h), tidal, inertial (29 h), high frequency (5 h) and residual currents  
  
Tides, eddies, internal wave surface expression
- **Current-Wave interactions:**  
  
current-induced shoaling, refraction
- **Wind-Wave** energy exchange

# Applications

- Nearshore modeling
- Engineering projects, ship navigation, vessel traffic control
- Mitigation of oil spills, ocean pollutants
- Shoreline protection, beach erosion
- Real-time sea state conditions for the tourism industry